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1 **Evaluation of short-term streamflow prediction methods in**

2 **urban river basins**

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15 **Abstract**

16 Efficient and accurate streamflow predictions are important for urban water management.
17 Data-driven models, especially neural network (NN) models can predict streamflow fast,
18 while the results are uncertain in some complex river systems. Physically based models
19 can reveal the underlying physics, but it is relatively slow and computationally costly. This

20 work focuses on evaluating the reliability of three NN models (artificial neural networks
21 (ANN), long short-term memory networks (LSTM), adaptive neuro-fuzzy inference
22 system (ANFIS)) and one physically based model (SOBEK) in terms of efficiency and
23 accuracy for average and peak streamflow simulation. All the models are applied for a tidal
24 river and a mountainous river in Shenzhen. The results show that, the ANN model
25 calculates fastest since the hidden layer's structure is simple. The LSTM model is reliable
26 in average streamflow simulation in tidal river with the lowest bias while the ANFIS model
27 has the best accuracy for peak streamflow simulation. Furthermore, the SOBEK model
28 shows reliability in simulating average and peak streamflow in mountainous river due to
29 its ability to capture uneven spatial rainfall in the area. Overall, the results indicate that the
30 LSTM model can be a helpful supplementary to physically based models in streamflow
31 simulation of complex urban river systems, by giving fast streamflow predictions with
32 usually acceptable accuracy. Our results can provide helpful information for hydrological
33 engineers in the application of flooding early warning and emergency preparedness in the
34 context of flooding risk management.

35 **Keywords:** streamflow simulation, neural network models, SOBEK model, urban rivers

36 **1. Introduction**

37 Urban flooding has become a threat to urban water security and has increased in frequency
38 in recent decades (Ziegler et al., 2012). China is prone to urban flooding, and many of
39 Chinese cities have been experiencing a large increases in urban flooding in recent years
40 (Duan et al., 2016). Between 2008 and 2010, 218 Chinese cities endured at least one urban
41 flooding event while more than 100 cities experienced three urban flooding events,

42 including major cities like Beijing, Shanghai, Guangzhou and Shenzhen (Jiang et al., 2018;
43 Song and Li, 2019). Tidal river basins and mountainous river basins are two typical urban
44 river basins with large impervious areas that are vulnerable to urban flooding (Archetti et
45 al., 2011; Davenport et al., 2004; Dawson et al., 2008). In most tidal river basins of China,
46 urban flooding is mainly caused by large amounts of rainfall-runoff at the same time of
47 sustained high tide at outlet, in which the rivers cannot convey large water volume during
48 high precipitation period (Lian et al., 2013; Orton et al., 2020). Urban flooding also occurs
49 in mountainous river basins due to its uneven slope distribution (Ballesteros-Cánovas et al.,
50 2015). It causes the streamflow to quickly move from high-altitude areas to low-altitude
51 areas, and impervious areas lead to slow streamflow infiltration, so a large amount of
52 streamflow will accumulate in a short time (Chen et al., 2008). Urban flooding can result
53 in disasters that cause enormous public and private property losses and casualties. In March
54 2014, Shenzhen experienced a 50-year rainfall event, paralyzing the urban sewer system
55 and surface water flows with more than 200 inundation areas (Xu et al., 2020). In 2016,
56 weeks of torrential rainfall during the monsoon season led to severe urban flooding, which
57 submerged 28 provinces and impacted 60 million people in China (Jiang et al., 2018).
58 Therefore, streamflow prediction in tidal river basins and mountainous river basins is
59 necessary and crucial for public safety management and social development.

60 Urban flooding is characterized by short duration and high intensity, making it difficult to
61 predict. The rapid forecasting and prediction of urban flooding can minimize potential
62 losses. Therefore, efficient and accurate simulations of the streamflow caused by urban
63 flooding is of great concern for urban water resource management and decision makers.
64 Hydrologists have paid increasing attention and efforts to developing urban flooding

65 models. Physically based models have been widely used for urban flooding prediction in
66 recent decades (Anghileri et al., 2016; Botto et al., 2018; Chen et al., 2016). They take the
67 dynamics of the hydrological cycle process into account, build a hydrodynamic equation
68 set based on the characteristics of runoff generation in the basin, and simulate the rainfall
69 runoff response. They can fully simulate the whole rainfall-runoff process of a river basin
70 (Kim and Mohanty, 2017). Physically based models, however, also present several
71 challenges. For instance, these models need large basic data for model set-up and
72 calibration (Li et al., 2020), and the quality of the simulation results depends on the quality
73 and availability of the input data (Yoon et al., 2011). Additionally, the calculation of
74 physically based models is relatively slow, and the computational cost is expensive due to
75 considerable data calibration and validation.

76 Due to the slow simulation process of physically based models (Yang et al., 2020), data-
77 driven models have gained considerable attention in hydrology in recent years due to their
78 rapid simulation capacity (Ahani et al., 2018). Data-driven models are therefore seen as
79 alternatives to physically based models. They consider the input and output data, without
80 using any of the physical processes (Wang and Yao, 2013). Among the various data-driven
81 models, neural network (NN) models are the most widely used techniques for streamflow
82 simulation and forecasting (Humphrey et al., 2016; Yang et al., 2020; Zhang et al., 2020).
83 Artificial neural network (ANN) model, can be seen as a black-box model, has been used
84 in river streamflow simulation because of its ability to mimic both linear, nonlinear and
85 hydrological systems (Aichouri et al., 2015; Kashani et al., 2016; Shoaib et al., 2014).
86 Among the data-driven models, it has a long development history, with the first studies
87 using the ANN model for streamflow prediction dating back to the early 1990s (Daniell,

88 1991; Halff et al., 1993). However, a drawback of ANN's hidden layer is that any
89 information about the sequential order of the inputs is lost. Long short-term memory
90 (LSTM) model overcomes the problem of ANN's hidden layer through a specially
91 designed architecture (Kratzert et al., 2018). This architecture has memory cells replacing
92 the traditional hidden layer. The memory cells could store, write and read data via gates
93 that open and close (Zhang et al., 2018). This can overcome the problem of the ANN model
94 of learning long-term dependencies representing, for example, storage effects within
95 hydrological catchments, which may play a significant role for hydrological process
96 (Kratzert et al., 2018). With the rapid development of data-driven modelling approaches,
97 there has been a shift from black-box models to semantic-based fuzzy systems in recent
98 years (Ang and Quek, 2005). Adaptive network-based fuzzy inference system (ANFIS)
99 model is one example of a semantic-based fuzzy system which conducts learning through
100 the minimization of global error within the model. However, there are several essential
101 limitations for NN models, such as the lack of explanation of the physical mechanism and
102 transparency of the simulation process and difficulties in explaining the results (Elshorbagy
103 et al., 2010).

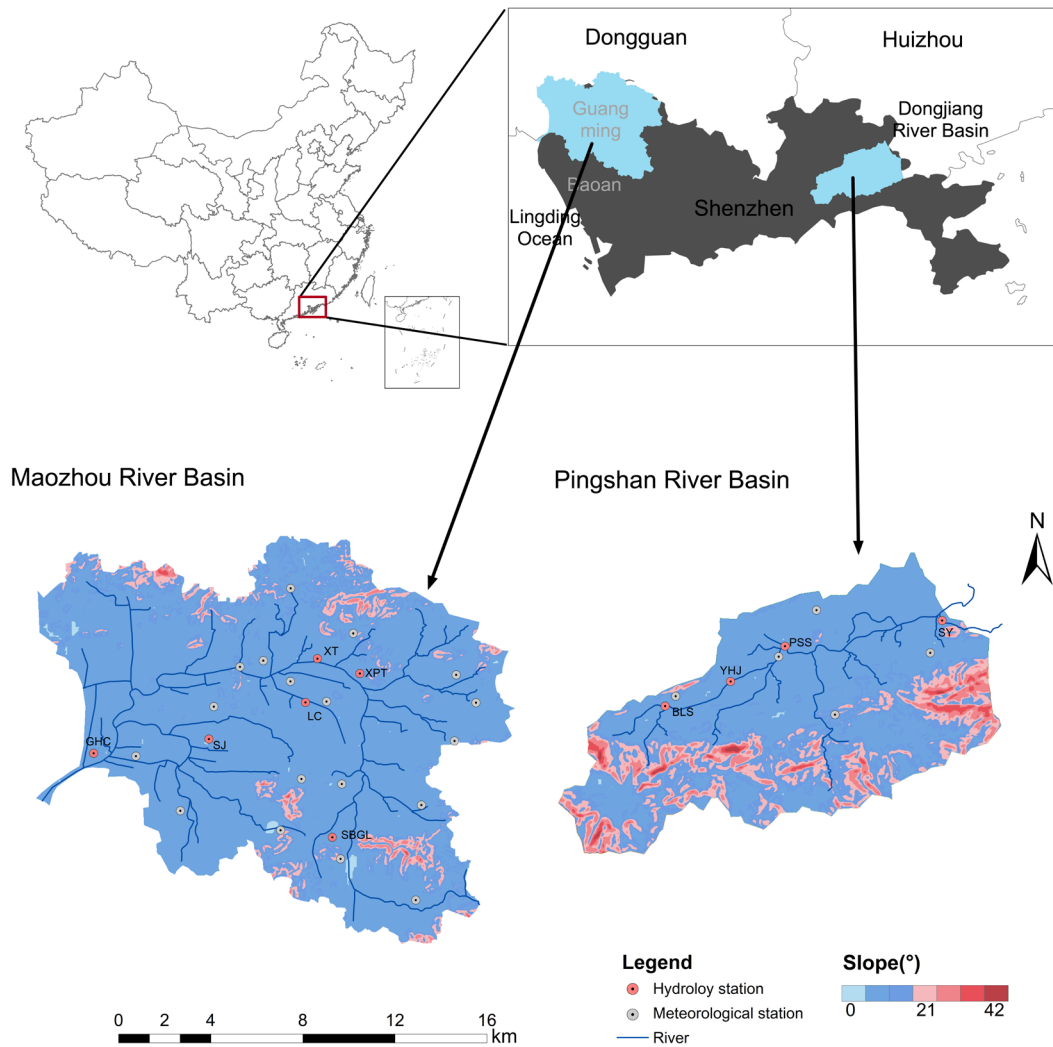
104 Through the research, we find that physically based models can make up for the
105 disadvantages of data-driven models that cannot simulate the physical process, and data-
106 driven models can make up for the slow simulation and prediction of physically based
107 models. However, due to the uncertainty of physical parameters, input parameters and basic
108 data, the simulation accuracy of the physically based model is also uncertain in some cases
109 (Hattermann et al., 2018; Her et al., 2019; Liu et al., 2017; Sikorska and Renard, 2017).
110 We therefore propose a hypothesis: for the streamflow simulation of a complex urban river

111 system (i.e., a tidal river basin and a mountainous river basin), physically based models
112 cannot fully generalize the physical process and the speed of the model is limited. Can
113 data-driven models work as the supplementary, help to better simulate the streamflow in
114 complex urban river systems? Existing studies that compared neural network models and
115 physically based models for streamflow simulation in urban areas have only concentrated
116 on data at the daily, monthly or annual scale (Chang and Chen, 2018; Mernild et al., 2018;
117 Nikpour et al., 2019; Schuol et al., 2008; Tikhamarine et al., 2020). There are very few
118 studies on the streamflow simulation based on hourly rainfall data. As the reliable
119 streamflow simulation and prediction plays a key role in confronting urban flooding risks,
120 a high temporal resolution precipitation dataset could have considerable influence on
121 model accuracy (Bruneau et al., 1995) and help deepen our understanding of the process
122 of streamflow, especially the process of extreme flooding events. The novelty of this study
123 is that it conducts an evaluation on the reliability of short-term streamflow prediction
124 methods driven by hourly rainfall data, with the goal to provide more suitable streamflow
125 simulation models for urban rivers.

126 This paper is organized as follows. Section 2 describes the study area, the data and method
127 used in this study. Section 3 shows the calibration and validation results of the models, and
128 the comparison of model performance in average and peak streamflow simulation. Section
129 4 and Section 5 present the discussion and conclusion, respectively.

130 **2. Materials**

131 **2.1 Study area**



132

133 Figure 1 Geography map and slope map (coordinate system: WGS-1984) of Maozhou
134 River Basin and Pingshan River Basin in Shenzhen city, China; the spatial distribution of
135 hydrological stations (red dots) and meteorological stations (white dots).

136 Shenzhen is a coastal city in China that is experiencing rapid economic development.

137 Urban flooding is one of the most devastating natural disasters that can occur in Shenzhen

138 (Ke et al., 2020). Due to its rapid urbanization, the frequency of urban flooding in Shenzhen

139 has increased in recent years (Shi et al., 2007; Yan et al., 2019). Maozhou River (MZR) is

140 a tidal river (Cui and Guo, 2006), located in the northwest of Shenzhen close to the borders
 141 of Dongguan, which is the largest watershed in Shenzhen (see Figure 1). It flows through
 142 Baoan District and Guangming District and finally flows into the Lingding Ocean.
 143 Pingshan River (PSR), a mountainous river with an average slope of 2.76% (Xiong et al.,
 144 2010), is located in north-eastern Shenzhen and close to the borders of Huizhou. The
 145 distribution of slopes in the two river basins is different (see Figure 1), resulting in different
 146 lag time period for the start of the rain to the peak of the flooding. Maozhou River can
 147 reach its peak in one hour, while Pingshan River can reach its peak in 40 minutes (SZN,
 148 2020). The comparison of general characteristics (including the river length, basin area,
 149 land use types, annual average rainfall, average slope and average elevation) of two rivers
 150 are shown in Table 1 (Chen et al., 2016; Cui and Guo, 2006; Peng et al., 2018; SMEEB,
 151 2018).

152 Table 1 General characteristics of two river basins (MZR and PSR).

River	Length (km)	Area (km ²)	Urban land	Forest land	Green land	Annual average rainfall(mm)	Average slope	Average elevation (m)
MZR	41.61	388.23	48%	32%	20%	1800	2.2‰	<25
PSR	22.14	129.40	40%	50%	10%	2073	2.76%	<82

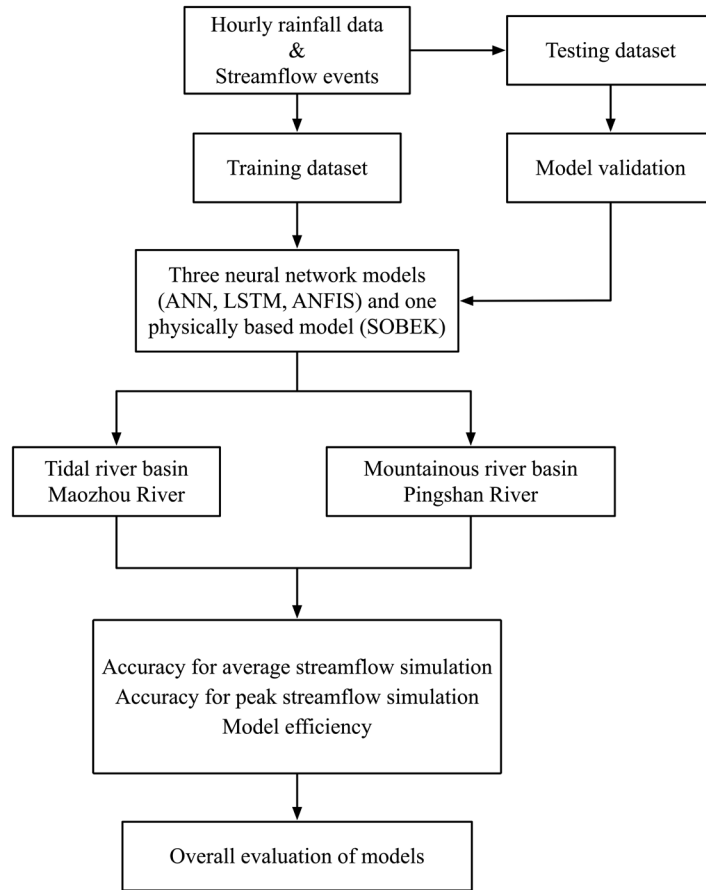
153 2.2 Data

154 In this study, digital elevation model (DEM) data, meteorological data, hourly streamflow
 155 observation data and river profiles were used. The DEM data is retrieved from the Shuttle
 156 Radar Topography Mission with a resolution of 30m (SRTM, 2020). Meteorological data,
 157 including hourly precipitation data, daily temperature (average, maximum and minimum),
 158 and daily average wind speed, are provided by the Meteorological Bureau of Shenzhen
 159 (SMB). Hourly streamflow observation data and the river profiles, including bed level,

160 channel slope, width of cross section and shape of cross section, are provided by the
161 Huadong Engineering Corporation Limited (ECIDI) and the Municipal Ecological
162 Environment Bureau of Shenzhen (SMEEB).

163 **2.3 Method**

164 As shown in Figure 2, we apply three neural network models (artificial neural networks
165 (ANN), long short-term memory networks (LSTM) and adaptive neuro-fuzzy inference
166 system (ANFIS)) with hourly rainfall data for streamflow simulation in a tidal river basin
167 (Maozhou River) and a mountainous river basin (Pingshan River) in Shenzhen, China, and
168 use a physically based model (the SOBEK model) as a reference. The introduction of the
169 four models can be found in the supplementary materials. We attempt to assess model
170 performance in three parts: model accuracy in average and peak streamflow simulation,
171 model accuracy. Meanwhile, we examine whether neural network models can make up for
172 physically based models in complex urban river systems in terms of accuracy.



173

174

Figure 2 The general framework of this study.

175 2.4 The definition of streamflow event

176 Adams et al. (1986) found that less than 60 minutes intervals between two streamflow

177 events, causes the division of rainfall to have greater impacts on the rainfall characteristic

178 parameters, which will not be conducive to the statistics of rainfall characteristics. If the

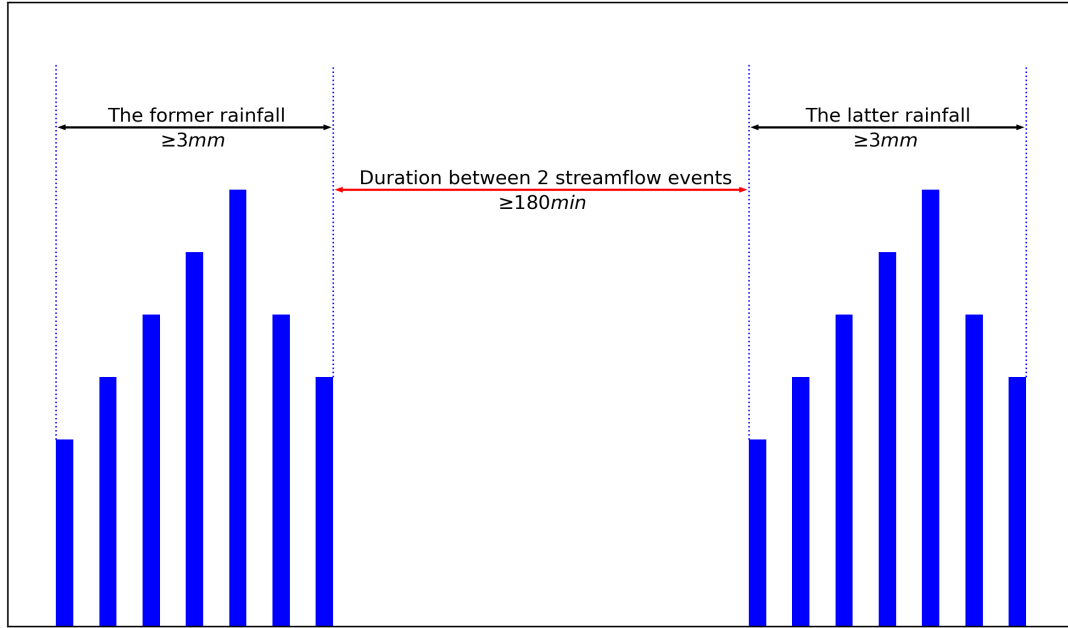
179 interval is set between 1-6 hours, the division of rainfall has a lower impact on rainfall

180 characteristic parameters and is more reasonable and scientific. Figure 3 shows the

181 definition of streamflow events in this paper. Considering the effects of rainfall confluence

182 time and rainfall duration, this study selects 180 minutes as the minimum interval between

183 two streamflow events, and the cumulative rainfall of each event is greater than 3 mm.



184

185

Figure 3 The definition of streamflow events in this research.

186

2.5 Assessment criteria

187

The Nash-Sutcliffe Efficiency (NSE) (Nash and Sutcliffe, 1970), the R square (R^2), the

188

percent bias (PBIAS) and the root-mean-square error (RMSE) are selected as assessment

189

criteria to evaluate the model results in calibration and validation periods. Meanwhile, the

190

Taylor diagram is applied to visualize the model performance in average streamflow

191

simulation (Taylor, 2001). In addition to the standard deviation and correlation coefficient

192

(r), the center root-mean-square errors (CRMSEs) are used in the Taylor diagram. The

193

equations for computing these objective functions are given as follows:

194

$$NSE = 1 - \frac{\sum_{i=1}^n (X_{i,Obs} - X_{i,Sim})^2}{\sum_{i=1}^n (X_{i,Obs} - \bar{X}_{Obs})^2} \quad (1)$$

195

196

$$R^2 = \left(\frac{\sum_{i=1}^n (X_{i,Obs} - \bar{X}_{Obs}) * (X_{i,Sim} - \bar{X}_{Sim})}{\sqrt{\sum_{i=1}^n (X_{i,Obs} - \bar{X}_{Obs})^2} * \sqrt{\sum_{i=1}^n (X_{i,Sim} - \bar{X}_{Sim})^2}} \right)^2 \quad (2)$$

197

198

$$PBIAS = \frac{\sum_{i=1}^n (X_{i,Sim} - X_{i,Obs})}{\sum_{i=1}^n (X_{i,Obs})} * 100 \quad (3)$$

199

200

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (X_{i,Obs} - X_{i,Sim})^2}{n}} \quad (4)$$

201

202

$$CRMSE = \sqrt{\frac{\sum_{i=1}^n ((X_{i,Obs} - \bar{X}_{Obs}) - (X_{i,Sim} - \bar{X}_{Sim}))^2}{n}} \quad (5)$$

203

where $X_{i,Obs}$ and $X_{i,Sim}$ are the i-th observation and simulation data respectively, \bar{X}_{Obs} and

204

\bar{X}_{Sim} are the mean values of the observation and simulation data respectively, and n is the

205

sample size.

206

The NSE can measure the goodness of fit, from 0 to 1, where a value approaching 1 means

207

the simulations are closer to the observations. The R^2 and r value are used to express the

208

correlation of simulation data and observation data directly, and a value approaching 1

209

indicates a perfect correlation. The PBIAS measures the average tendency of the simulated

210

values to be larger or smaller than their observed ones. The optimal value of PBIAS is 0.0,

211

positive values indicate overestimation bias, whereas negative values indicate model

212

underestimation bias (Moriasi et al., 2007). The RMSE is used to evaluate how closely the

213

simulation data match the observation data and the values can range from 0 to $+\infty$ based

214

on the range of the data. The CRMSE is similar to the RMSE but has easier visualization

215

characteristics in Taylor diagram. It ranges from 0 to 1, where a value approaching 0 means

216

good simulation result.

217 **3. Results**

218 **3.1 Model accuracy in average streamflow simulation**

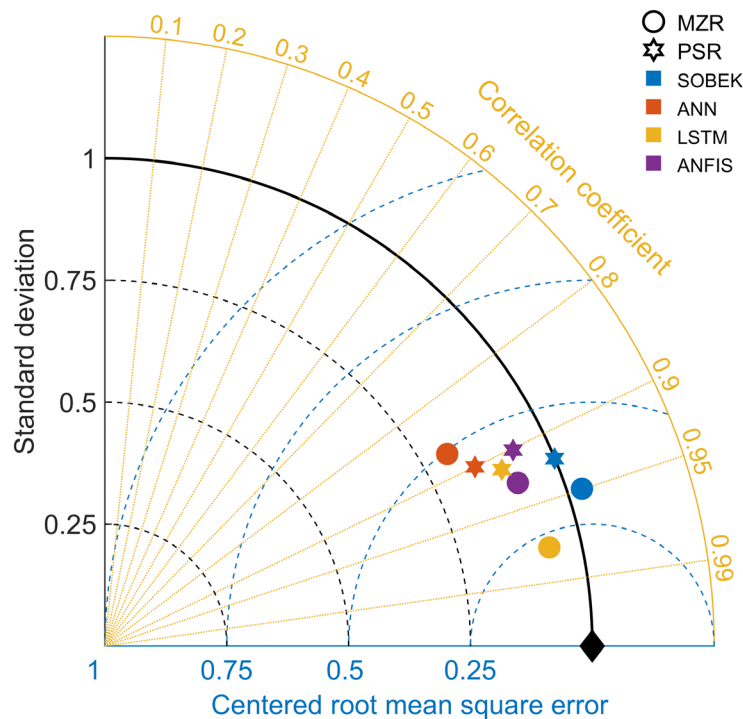
219 We first selected 59 streamflow events (49 events for calibration; 10 events for validation)
 220 for Maozhou River and 21 events (11 events for calibration; 10 events for validation) for
 221 Pingshan River. As seen from Table 2, four models can simulate the streamflow of two
 222 river basins satisfactorily with high NSE (larger than 0.80) and low absolute PBIAS value
 223 (less than 15%) and the ANN model is the best one among four models in calibration period
 224 in Maozhou River, the SOBEK model is the best in Pingshan River. Moreover, it can be
 225 seen that in the validation period, the LSTM model is the best in Maozhou River, and the
 226 SOBEK model is the best in Pingshan River. The details of the calibration and validation
 227 results can be found in appendices (Figure A1).

228 Table 2 Comparison of assessment criteria of four models for streamflow simulation
 229 in calibration and validation periods.

Basins	Models	NSE	R ²	RMSE (m ³ /s)	PBIAS (%)
MZR	Calibration				
	SOBEK	0.849	0.868	10.137	-12.029
	ANN	0.954	0.956	5.601	-0.850
	LSTM	0.922	0.939	7.284	-14.317
	ANFIS	0.852	0.854	10.029	4.071
	Validation				
	SOBEK	0.948	0.950	11.387	-0.542
	ANN	0.587	0.687	32.106	-17.860
	LSTM	0.976	0.977	7.725	-1.379
	ANFIS	0.883	0.884	17.123	0.121
PSR	Calibration				
	SOBEK	0.925	0.926	1.922	-6.760
	ANN	0.874	0.879	2.498	-0.702
	LSTM	0.888	0.892	2.358	6.461
	ANFIS	0.817	0.820	3.010	3.989
	Validation				
SOBEK	0.891	0.897	2.109	1.670	

ANN	0.645	0.646	3.805	6.480
LSTM	0.830	0.848	2.633	9.791
ANFIS	0.878	0.880	2.235	4.982

230 Visualizing the average streamflow simulation statistic results of standard deviation,
 231 CRMSE, and the correlation coefficient in the Taylor diagram (Figure 4) verifies the
 232 distinguish performance of each model. Generally, the LSTM model in Maozhou River
 233 (yellow dot) and the SOBEK model in Pingshan River (blue star) lead the reliability of
 234 average streamflow simulation regarding the statistical performance, characterized by
 235 relatively small standard deviation and CRMSE, and relatively higher r values. The details
 236 of the comparison between the observed and simulated results for the four models in terms
 237 of average streamflow in two river basins can be found in appendices (Figure A2 and A3).



238

239 Figure 4 Taylor diagram of the four models in two basins, with shapes and colours
 240 indicating simulation and observation data.

241 **3.2 Evaluation of model efficiency**

242 In Section 3.1, we evaluated the model accuracy in average streamflow simulation during
243 calibration and validation period. In this section, the model efficiency will be evaluated as
244 the computation time of models would be of great importance for the streamflow
245 simulation during short period.

246 The computation time of the models is related to the computer's CPU and memory. To
247 eliminate the impact of different computer configurations, the computer configuration we
248 adopted is as follows: i7-8700 CPU, 32G memory. Under the same computer configuration,
249 the computation time of the SOBEK model depends on the complexity of the constructed
250 river network and the amount of input rainfall data. As the river network of Maozhou River
251 is more complicated than Pingshan River, the SOBEK model validates for almost 1 hour
252 in Maozhou River and 30 minutes in Pingshan River. Unlike the physically based models,
253 rainfall data is the only input required in the neural network model (after calibration) to
254 obtain the prediction streamflow. Therefore, the computation time is much faster.

255 We also measured the validation time of the three neural network models. The time series
256 of the two basins for validation is approximately 350 hours. The ANN model is the quickest
257 one (the simulation time is 6 min for the MZR and 5 min for the PSR), and the LSTM
258 model is the slowest among three neural network models we used (12 min for the MZR
259 and 10 min for the PSR). The ANN model and LSTM model provided faster simulation
260 for Pingshan River than Maozhou River as there were fewer streamflow events in Pingshan
261 River. Fewer streamflow events mean less simulation time. The simulation time of the
262 ANFIS model was between the ANN model and the LSTM model, and there was no

263 difference in the simulation time of the ANFIS model in the two basins (approximately 8
264 min). Overall, it shows that the validation time of the neural network models is faster than
265 that of physically based model.

266 For the prediction time, we performed a rough test to estimate simulation time based on
267 the existing validation time. In Shenzhen, the streamflow from rainfall to peak value does
268 not exceed three hours (SZN, 2020). Therefore, we used a three-hour rainfall data for
269 testing. As seen in Table 3, under a future three-hour rainfall event, the SOBEK model has
270 a better performance in prediction time for each basin than in validation because the amount
271 of input rainfall data is smaller. The neural network models can increase a factor of 10-60
272 over the SOBEK model for 3-hours rainfall duration (the ANN model takes approximately
273 30 seconds, the LSTM model approximately 2 minutes and the ANFIS model
274 approximately 1 minute).

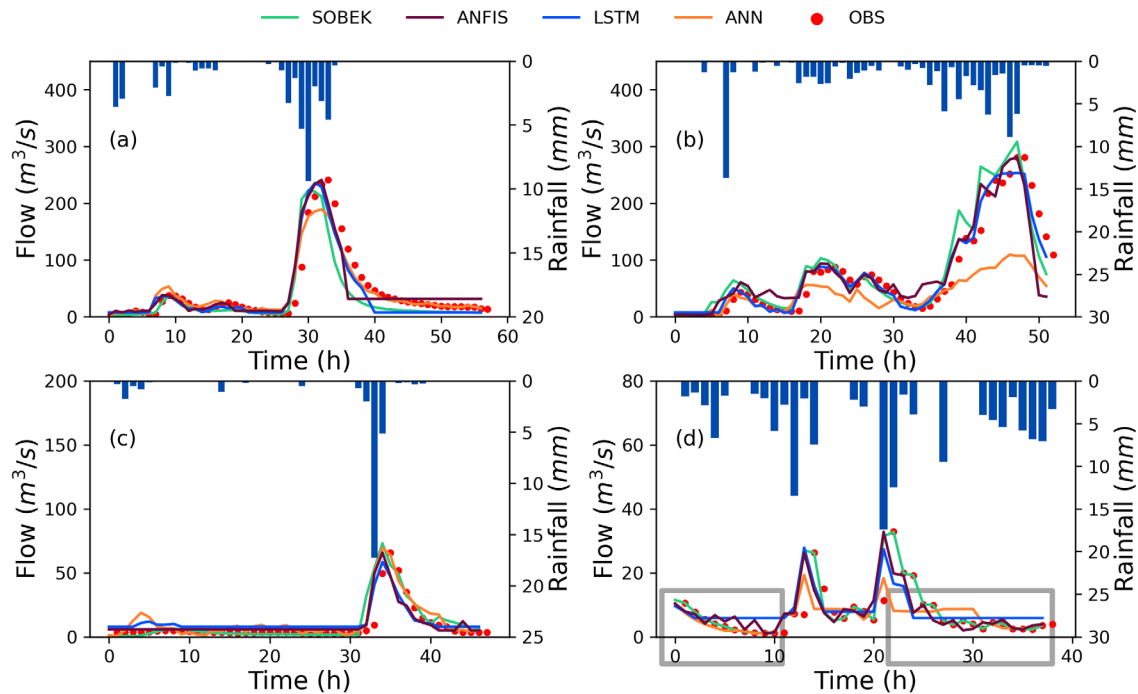
275 Table 3 The prediction time of 3-hour rainfall data using four models.

Model \ River	SOBEK	ANN	LSTM	ANFIS
MZR	30 min	≈30s	≈2 min	≈1 min
PSR	15min	≈30s	≈1.5 min	≈1 min

276 3.3 Model performance in predicting flooding events

277 To analyze the simulation performance of the four models during a flooding event process,
278 we compare the observation and simulation hydrographs of four flooding events (the
279 rainfall characteristics of these events are short duration and high intensity) using the four
280 models in the two river basins ((a) is SJ station, (b) is LC station and (c) is XT station in
281 MZR, (d) is PSS station in PSR see figure 1) in Figure 5. The ANFIS model is the best
282 among the four models for the simulation of peak streamflow during a flooding event,

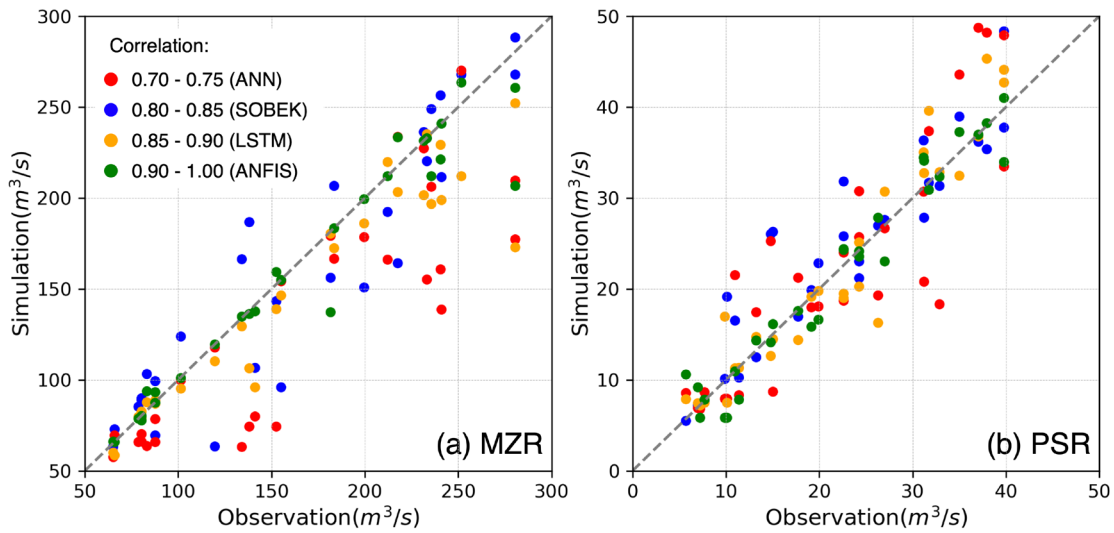
283 while the ANN model is the worst, which relies in the simulated values of the ANN model
 284 fluctuate abnormally compared with the observed values in large volume streamflow
 285 (especially at LC station see Figure 5(b)). The LSTM model is not very effective in
 286 simulating small volume flow values in flooding events (the LSTM model cannot reflect
 287 the fluctuating state of small volume streamflow, see Figure 5 (d), highlighted by a grey
 288 box). The SOBEK model shows abnormal fluctuate values between the time of 40-50 hours
 289 in Figure 5(b) (see green line).



290
 291 Figure 5 Observed and simulated hydrographs of streamflow in the two basins.
 292 (a) - SJ station, (b) -LC station (c) - XT station in MZR, (d) - PSS station in PSR.

293 In addition, a comparison study on the correlation of observation data and simulation
 294 results is conducted, and the result is shown in Figure 6. As seen in Figure 6, the four
 295 models all have a high correlation (higher than 0.70) between the observation data and the
 296 simulation values in two river basins. Compared with other models, the distribution of the

297 ANFIS model (green dots) is more concentrated in Maozhou River (Figure 6(a)), and the
 298 SOBEK model (blue dots) is more concentrated in Pingshan River (Figure 6(b)),
 299 respectively. This reveals that the ANFIS and SOBEK model have great potential to
 300 simulate peak streamflow well in tidal river basin and in mountainous river basin,
 301 respectively.



302
 303 Figure 6 Observed and simulated hydrographs of streamflow in MZR (a) and PSR (b)
 304 basins.

305 4. Discussion

306 From the above research, all four models have potential to simulate the streamflow of urban
 307 river basins well. Table 4 ranks the four models' performance in terms of accuracy for
 308 average streamflow simulation (AAS), accuracy for peak streamflow simulation (APS),
 309 model efficiency and overall choice in two river basins.

310 Table 4 Rank performance of four models in two river basins.

	Tidal river basin				Mountainous river basin			
	SOBEK	ANN	LSTM	ANFIS	SOBEK	ANN	LSTM	ANFIS
AAS	2 nd	4 th	1 st	3 rd	1 st	4 th	2 nd	3 rd
APS	3 rd	4 th	2 nd	1 st	1 st	4 th	3 rd	2 nd
Model	4 th	1 st	2 nd	3 rd	4 th	1 st	2 nd	3 rd
Efficiency								
Overall			√		√			
Choice								

311 Of the four models, the LSTM model shows a good ability in simulating average
312 streamflow well in tidal river basin. The internal memory cells of the LSTM model
313 ('forgotten gate' and 'memory gate') have the ability to filter data and memory data
314 features making as neural network functions to simulate the average streamflow process
315 could be the reason for its good performance (Kratzert et al., 2018; Sahoo et al., 2019;
316 Sudriani et al., 2019). Besides, the low streamflow has little impact on the average
317 streamflow prediction accuracy in tidal river basin and the physical-based models
318 sometimes cannot apply well in tidal basins due to some uncertainty sources, such as
319 excessive rainfall data and the tidal effects (Jung et al., 2018), so the LSTM model can
320 simulate better than the SOBEK model.

321 The uneven slope distribution has great influence on the average streamflow simulation
322 accuracy in mountainous river basin. The SOBEK model can therefore, exhibit better due

323 to the ability of physically based models to response to the rainfall-streamflow process
324 (Noor et al., 2014). However, the ANN model does not show high reliability in validation
325 period. According to Hu et al. (2018) and Ahmad and Simonovic (2005), the hidden layer
326 function of the ANN model has limitations and challenges when simulate insufficient data.
327 The input streamflow event of validation period is not enough, so the ANN model cannot
328 validate well. Moreover, the results of the ANFIS model are similar to those of the LSTM
329 model and more stable than those of the ANN model, which relies on the fact that the
330 ANFIS model combines the relationship structure of neural network models with the
331 decision-making mechanism of fuzzy logic (Amutha and Porchelvan, 2011; Vetrivel and
332 Elangovan, 2017).

333 Compared to physically based model, all neural network models have the ability to
334 simulate and predict quickly. The physically based model needs several data for model set-
335 up and the calculation time depends on the complexity of urban river system. Therefore,
336 physically based model has disadvantages in prediction time (Ke et al., 2020; Sun et al.,
337 2017). The prediction speed of the ANN model is the best among the three neural network
338 models we used. The ANN model can complete nonlinear predictions by adjusting the
339 number and type of neurons in the hidden layer and the weights carried by each neuron
340 (Navale and Singh, 2020). The model structure is relatively simple and can verify a large
341 amount of data, so the prediction time is the shortest. Compared to the ANN model, the
342 LSTM model requires more data for training and validation, so the calculation time is
343 longer than that of the ANN model. The ANFIS model combines the characteristics of
344 fuzzy systems and neural networks and adjusts the model by adjusting the type and number
345 of membership functions. Due to the complex structure and gradient learning, the

346 computational cost of ANFIS is very high, and it has more difficulty dealing with a large
347 amount of input data (Salleh et al., 2017).

348 For peak streamflow prediction, the ANN model performs the largest underestimations in
349 two river basins. However, it exhibits good performance in low volume streamflow
350 simulation. The poor performance of the ANN model on peak streamflow simulation is in
351 line with Sudheer et al. (2003) who highlighted that the ANN model tends to underestimate
352 the peak streamflow even after data transformation and the learning process of an ANN
353 model will reward a correct response of the system to input by increasing the strength of
354 the current matrix of nodal weights. Likewise, the LSTM model is unable to simulate the
355 small volume streamflow in a flood event well in mountainous river basin (see Figure 5
356 (d)), as having a continual value of streamflow for a high quantile of training data seems
357 to pose difficulty for the LSTM model to learn and calibrate (Kratzert et al., 2018). The
358 ANFIS model has the best accuracy in tidal river basin and the SOBEK model has the best
359 accuracy in mountainous river basin, respectively. The peak streamflow value in tidal river
360 basin is relatively large, the ANFIS model has a greater ability to train large data, as it
361 combines the characteristics of fuzzy systems and neural networks, making it most
362 appropriate for tidal river basins. The complex mechanisms of river system and lack of
363 data are two main challenges in peak streamflow simulation of mountainous river basins
364 (Zhang et al., 2013). The physically based model can reflect the complex mechanisms
365 between rainfall and streamflow, and neural network models may have insufficient training
366 data due to the low frequency of peak streamflow in mountainous river basins (Sudheer et
367 al., 2003; Yang et al., 2019).

368 **5. Conclusion**

369 An evaluation on the performance of three neural network models and one physically based
370 model for the streamflow simulation driven by hourly precipitation data in one tidal river
371 basin and one mountainous river basin of Shenzhen has been conducted in this study. The
372 following major findings are drawn:

373 (1) The four models we used are able to capture streamflow simulation well in two river
374 basins. Specifically, the LSTM model is reliable in terms of average streamflow simulation
375 in tidal river basin with the lowest bias but underestimates the small volume streamflow. It
376 is suitable for tidal river basins that low streamflow has little impact on the simulation
377 accuracy. The SOBEK model shows reliability in simulating average streamflow in
378 mountainous river basin, while needs large basic data for model calibration and validation.

379 (2) All neural network models used in this research present high simulation speed. The
380 three neural network models can predict a three-hour rainfall event in less than 2 minutes.
381 The ANN model shows great reliability in prediction speed, is suitable for scenarios that
382 require high forecasting speed such as emergency flooding management.

383 (3) The ANFIS model has best accuracy for peak streamflow simulation in tidal river basin,
384 is suitable for extreme flooding prediction. For example, in urban flooding management,
385 decision makers need to obtain extreme value of a flooding event in order to facilitate
386 flooding management and decision-making. The SOBEK model has best accuracy of peak
387 streamflow simulation in mountainous river basin, is suitable for river basins with complex
388 mechanism systems.

389 (4) Overall, the LSTM model can compensate for physically based models in streamflow
390 simulation in complex urban river systems, by giving fast streamflow predictions with
391 usually acceptable accuracy.

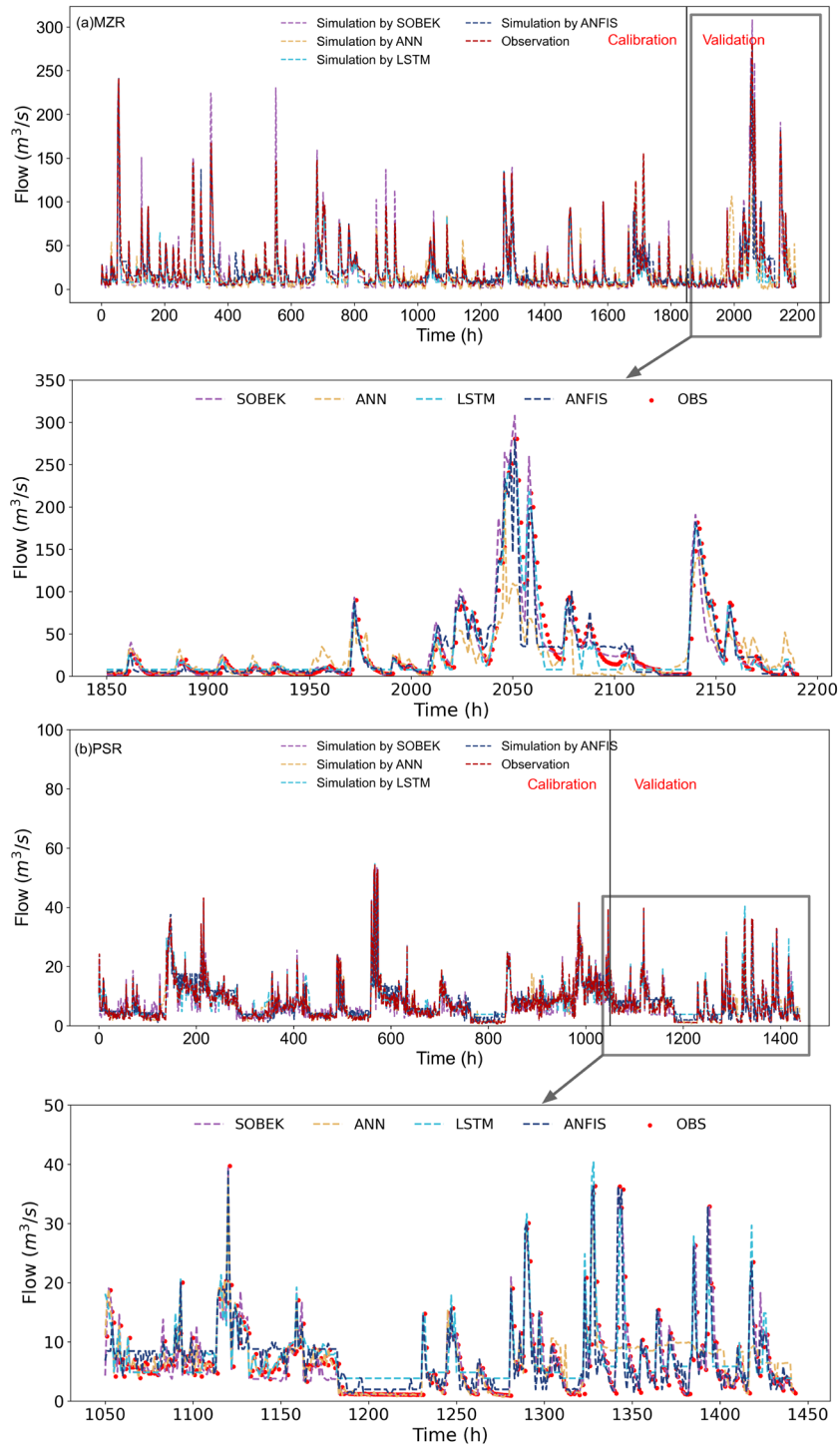
392 Our research provides scientific support for the application in flood early warning and
393 emergency preparedness in the context of flood risk management in the urban area.

394 **Declaration of Competing Interest**

395 The authors declare that they have no known competing financial interests or personal
396 relationships that could have appeared to influence the work reported in this paper.

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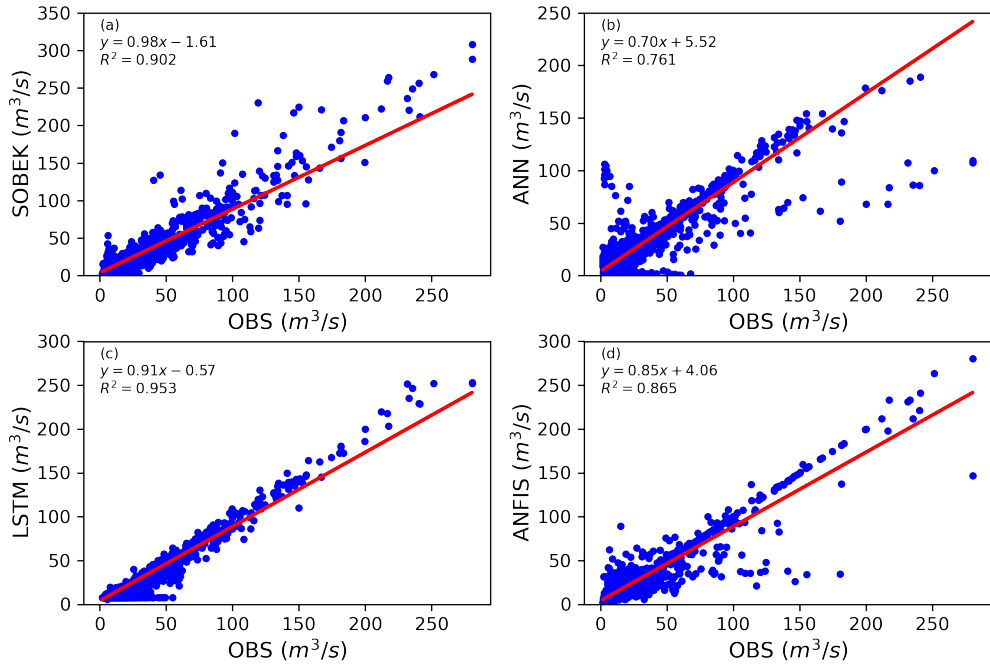


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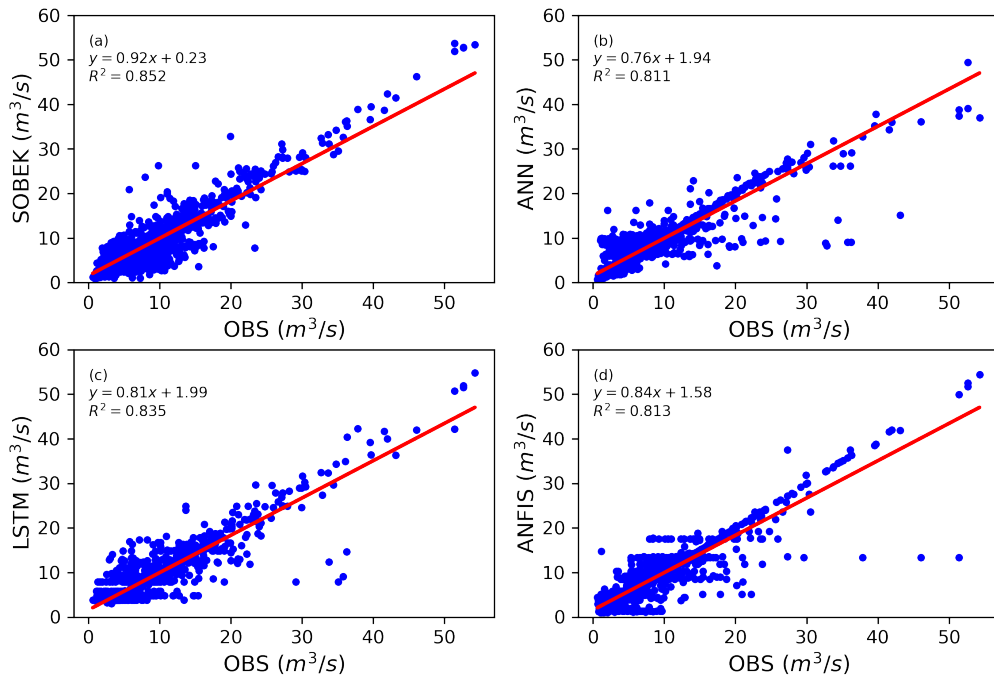
408

Figure A1 Details of calibration and validation results in two basins (a-MZR, b-PSR).



409

410 Figure A2 Comparison of observation (x-axis) and simulated results (y-axis) for the four
 411 models in terms of streamflow in MZR.



412

413 Figure A3 Comparison of observation (x-axis) and simulated results (y-axis) for the four
 414 models in terms of streamflow in PSR.

416 **Reference**

- 417 Adams, B.J., Fraser, H.G., Howard, C.D.D., Sami Hanafy, M., 1986. Meteorological data
418 analysis for drainage system design. *J. Environ. Eng.* 112, 827–848.
- 419 Ahani, A., Shourian, M., Rahimi Rad, P., 2018. Performance Assessment of the Linear,
420 Nonlinear and Nonparametric Data Driven Models in River Flow Forecasting.
421 *Water Resour. Manag.* 32, 383–399. <https://doi.org/10.1007/s11269-017-1792-5>
- 422 Ahmad, S., Simonovic, S.P., 2005. An artificial neural network model for generating
423 hydrograph from hydro-meteorological parameters. *J. Hydrol.* 315, 236–251.
424 <https://doi.org/10.1016/j.jhydrol.2005.03.032>
- 425 Aichouri, I., Hani, A., Bougherira, N., Djabri, L., Chaffai, H., Lallahem, S., 2015. River
426 Flow Model Using Artificial Neural Networks. *Energy Procedia* 74, 1007–1014.
427 <https://doi.org/10.1016/j.egypro.2015.07.832>
- 428 Amutha, R., Porchelvan, P., 2011. Seasonal Prediction of Groundwater Levels Using
429 Anfis 1, 98–108.
- 430 Ang, K.K., Quek, C., 2005. Rspop: Rough set-based pseudo outer-product fuzzy rule
431 identification algorithm. *Neural Comput.* 17, 205–243.
- 432 Anghileri, D., Giudici, F., Castelletti, A., Burlando, P., 2016. Advancing reservoir
433 operation description in physically based hydrological models, in: EGUGA. pp.
434 EPSC2016-10097.
- 435 Archetti, R., Bolognesi, A., Casadio, A., Maglionico, M., 2011. Development of flood
436 probability charts for urban drainage network in coastal areas through a simplified
437 joint assessment approach. *Hydrol. Earth Syst. Sci.* 15, 3115–3122.
438 <https://doi.org/10.5194/hess-15-3115-2011>
- 439 Ballesteros-Cánovas, J.A., Rodríguez-Morata, C., Garófano-Gómez, V., Rubiales, J.M.,
440 Sánchez-Salguero, R., Stoffel, M., 2015. Unravelling past flash flood activity in a
441 forested mountain catchment of the Spanish Central System. *J. Hydrol.* 529, 468–
442 479. <https://doi.org/10.1016/j.jhydrol.2014.11.027>
- 443 Botto, A., Belluco, E., Camporese, M., 2018. Multi-source data assimilation for
444 physically based hydrological modeling of an experimental hillslope. *Hydrol. Earth*
445 *Syst. Sci.* 22, 4251.
- 446 Bruneau, P., Gascuel-Oudou, C., Robin, P., Merot, P., Beven, K., 1995. Sensitivity to
447 space and time resolution of a hydrological model using digital elevation data.
448 *Hydrol. Process.* 9, 69–81. <https://doi.org/10.1002/hyp.3360090107>
- 449 Chang, W., Chen, X., 2018. Monthly rainfall-runoff modeling at watershed scale: A
450 comparative study of data-driven and theory-driven approaches. *Water*
451 (Switzerland) 10, 1–21. <https://doi.org/10.3390/w10091116>

- 452 Chen, R.S., Lu, S.H., Kang, E.S., Ji, X. Bin, Zhang, Z., Yang, Y., Qing, W., 2008. A
453 distributed water-heat coupled model for mountainous watershed of an inland river
454 basin of Northwest China (I) model structure and equations. *Environ. Geol.* 53,
455 1299–1309. <https://doi.org/10.1007/s00254-007-0738-2>
- 456 Chen, Y., Li, J., Xu, H., 2016. Improving flood forecasting capability of physically based
457 distributed hydrological models by parameter optimization. *Hydrol. Earth Syst. Sci.*
458 20, 375.
- 459 Cui, X., Guo, R., 2006. Hydrological characteristics of Maozhou River Basin. *China*
460 *Rural Water Hydropower* 9, 57–60.
- 461 Daniell, T.M., 1991. Neural networks. Applications in hydrology and water resources
462 engineering, in: National Conference Publication- Institute of Engineers. Australia.
- 463 Davenport, A.J., Gurnell, A.M., Armitage, P.D., 2004. Habitat survey and classification
464 of urban rivers. *River Res. Appl.* 20, 687–704. <https://doi.org/10.1002/rra.785>
- 465 Dawson, R.J., Speight, L., Hall, J.W., Djordjevic, S., Savic, D., Leandro, J., 2008.
466 Attribution of flood risk in urban areas. *J. Hydroinformatics* 10, 275–288.
467 <https://doi.org/10.2166/hydro.2008.054>
- 468 Duan, W., He, B., Nover, D., Fan, J., Yang, G., Chen, W., Meng, H., Liu, C., 2016.
469 Floods and associated socioeconomic damages in China over the last century. *Nat.*
470 *Hazards* 82, 401–413. <https://doi.org/10.1007/s11069-016-2207-2>
- 471 Elshorbagy, A., Corzo, G., Srinivasulu, S., Solomatine, D.P., 2010. Experimental
472 investigation of the predictive capabilities of data driven modeling techniques in
473 hydrology - Part 1: Concepts and methodology. *Hydrol. Earth Syst. Sci.* 14, 1931–
474 1941. <https://doi.org/10.5194/hess-14-1931-2010>
- 475 Halff, A.H., Halff, H.M., Azmoodeh, M., 1993. Predicting runoff from rainfall using
476 neural networks, in: *Engineering Hydrology*. ASCE, pp. 760–765.
- 477 Hattermann, F.F., Vetter, T., Breuer, L., Su, B., Daggupati, P., Donnelly, C., Fekete, B.,
478 Florke, F., Gosling, S.N., Hoffmann, P., Liersch, S., Masaki, Y., Motovilov, Y.,
479 Muller, C., Samaniego, L., Stacke, T., Wada, Y., Yang, T., Krysnova, V., 2018.
480 Sources of uncertainty in hydrological climate impact assessment: A cross-scale
481 study. *Environ. Res. Lett.* 13. <https://doi.org/10.1088/1748-9326/aa9938>
- 482 Her, Y., Yoo, S.H., Cho, J., Hwang, S., Jeong, J., Seong, C., 2019. Uncertainty in
483 hydrological analysis of climate change: multi-parameter vs. multi-GCM ensemble
484 predictions. *Sci. Rep.* 9, 1–22. <https://doi.org/10.1038/s41598-019-41334-7>
- 485 Hu, C., Wu, Q., Li, H., Jian, S., Li, N., Lou, Z., 2018. Deep learning with a long short-
486 term memory networks approach for rainfall-runoff simulation. *Water (Switzerland)*
487 10, 1–16. <https://doi.org/10.3390/w10111543>
- 488 Humphrey, G.B., Gibbs, M.S., Dandy, G.C., Maier, H.R., 2016. A hybrid approach to
489 monthly streamflow forecasting: Integrating hydrological model outputs into a
490 Bayesian artificial neural network. *J. Hydrol.* 540, 623–640.
491 <https://doi.org/10.1016/j.jhydrol.2016.06.026>

- 492 Jiang, Y., Zevenbergen, C., Ma, Y., 2018. Urban pluvial flooding and stormwater
493 management: A contemporary review of China's challenges and "sponge cities"
494 strategy. *Environ. Sci. Policy* 80, 132–143.
495 <https://doi.org/10.1016/j.envsci.2017.11.016>
- 496 Jung, Sungho • Cho, Hyoseob • Kim, Jeongyup • Lee, Giha aD, D.P.E.E., bW R H R F
497 C O, K.N.U., E, M., 2018. Prediction of water level in a tidal river using a deep-
498 learning based LSTM model 51, 1207–1216.
499 <https://doi.org/10.3741/JKWRA.2018.51.12.1207>
- 500 Kashani, M., Ghorbani, M.A., Dinpashoh, Y., Shahmorad, S., 2016. Integration of
501 Volterra model with artificial neural networks for rainfall-runoff simulation in
502 forested catchment of northern Iran. *J. Hydrol.* 540, 340–354.
503 <https://doi.org/10.1016/j.jhydrol.2016.06.028>
- 504 Ke, Q., Tian, X., Bricker, J., Tian, Z., Guan, G., Cai, H., Huang, X., Yang, H., Liu, J.,
505 2020. Urban pluvial flooding prediction by machine learning approaches – a case
506 study of Shenzhen city, China. *Adv. Water Resour.* 145, 103719.
507 <https://doi.org/10.1016/j.advwatres.2020.103719>
- 508 Kim, J., Mohanty, B.P., 2017. A physically based hydrological connectivity algorithm for
509 describing spatial patterns of soil moisture in the unsaturated zone. *J. Geophys. Res.*
510 122, 2096–2114. <https://doi.org/10.1002/2016JD025591>
- 511 Kratzert, F., Klotz, D., Brenner, C., Schulz, K., Herrnegger, M., 2018. Rainfall – runoff
512 modelling using Long Short-Term Memory (LSTM) networks 6005–6022.
- 513 Li, P., Zha, Y., Shi, L., Tso, C.H.M., Zhang, Y., Zeng, W., 2020. Comparison of the use
514 of a physical-based model with data assimilation and machine learning methods for
515 simulating soil water dynamics. *J. Hydrol.* 584, 124692.
516 <https://doi.org/10.1016/j.jhydrol.2020.124692>
- 517 Lian, J.J., Xu, K., Ma, C., 2013. Joint impact of rainfall and tidal level on flood risk in a
518 coastal city with a complex river network: A case study of Fuzhou City, China.
519 *Hydrol. Earth Syst. Sci.* 17, 679–689. <https://doi.org/10.5194/hess-17-679-2013>
- 520 Liu, Y.R., Li, Y.P., Huang, G.H., Zhang, J.L., Fan, Y.R., 2017. A Bayesian-based
521 multilevel factorial analysis method for analyzing parameter uncertainty of
522 hydrological model. *J. Hydrol.* 553, 750–762.
523 <https://doi.org/10.1016/j.jhydrol.2017.08.048>
- 524 Mernild, S.H., Liston, G.E., Hiemstra, C.A., Yde, J.C., Casassa, G., 2018. Annual river
525 runoff variations and trends for the Andes Cordillera. *J. Hydrometeorol.* 19, 1167–
526 1189. <https://doi.org/10.1175/JHM-D-17-0094.1>
- 527 Moriasi, D.N., Arnold, J.G., Van Liew, M.W., Bingner, R.L., Harmel, R.D., Veith, T.L.,
528 2007. Model evaluation guidelines for systematic quantification of accuracy in
529 watershed simulations. *Trans. ASABE* 50, 885–900.
- 530 Nash, J.E., Sutcliffe, J. V., 1970. River flow forecasting through conceptual models part I
531 - A discussion of principles. *J. Hydrol.* 10, 282–290. <https://doi.org/10.1016/0022->

532 1694(70)90255-6

533 Navale, A., Singh, C., 2020. Topographic sensitivity of WRF-simulated rainfall patterns
534 over the North West Himalayan region. *Atmos. Res.* 242, 105003.
535 <https://doi.org/10.1016/j.atmosres.2020.105003>

536 Nikpour, M.R., Sanikhani, H., Babelan, S.M., Amuqin, S.N., 2019. Archive of SID Daily
537 Rainfall – Runoff Modeling of Darreh-Rud River in Ardabil Province , Iran Archive
538 of SID. pp. 144–146.

539 Noor, H., Vafakhah, M., Taheriyoun, M., Moghadasi, M., 2014. Hydrology modelling in
540 Taleghan mountainous watershed using SWAT. *J. Water L. Dev.* 20, 11–18.
541 <https://doi.org/10.2478/jwld-2014-0003>

542 Orton, P.M., Conticello, F.R., Cioffi, F., Hall, T.M., Georgas, N., Lall, U., Blumberg,
543 A.F., MacManus, K., 2020. Flood hazard assessment from storm tides, rain and sea
544 level rise for a tidal river estuary. *Nat. Hazards* 102, 729–757.
545 <https://doi.org/10.1007/s11069-018-3251-x>

546 Peng, J., Wei, H., Wu, W., Liu, Y., Wang, Y., 2018. Storm flood disaster risk assessment
547 in urban area based on the simulation of land use scenarios: A case of Maozhou
548 Watershed in Shenzhen City. *ACTA Ecol. Sin.* 38, 3741–3755.

549 Sahoo, B.B., Jha, R., Singh, A., Kumar, D., 2019. Long short-term memory (LSTM)
550 recurrent neural network for low-flow hydrological time series forecasting. *Acta*
551 *Geophys.* 67, 1471–1481. <https://doi.org/10.1007/s11600-019-00330-1>

552 Salleh, M.N.M., Talpur, N., Hussain, K., 2017. Adaptive neuro-fuzzy inference system:
553 Overview, strengths, limitations, and solutions, in: *International Conference on Data*
554 *Mining and Big Data*. Springer, pp. 527–535.

555 Schuol, J., Abbaspour, K.C., Yang, H., Srinivasan, R., Zehnder, A.J.B., 2008. Modeling
556 blue and green water availability in Africa. *Water Resour. Res.* 44, 1–18.
557 <https://doi.org/10.1029/2007WR006609>

558 Shi, P.J., Yuan, Y., Zheng, J., Wang, J.A., Ge, Y., Qiu, G.Y., 2007. The effect of land
559 use/cover change on surface runoff in Shenzhen region, China. *Catena* 69, 31–35.
560 <https://doi.org/10.1016/j.catena.2006.04.015>

561 Shoaib, M., Shamseldin, A.Y., Melville, B.W., 2014. Comparative study of different
562 wavelet based neural network models for rainfall-runoff modeling. *J. Hydrol.* 515,
563 47–58. <https://doi.org/10.1016/j.jhydrol.2014.04.055>

564 Sikorska, A.E., Renard, B., 2017. Calibrating a hydrological model in stage space to
565 account for rating curve uncertainties: general framework and key challenges. *Adv.*
566 *Water Resour.* 105, 51–66. <https://doi.org/10.1016/j.advwatres.2017.04.011>

567 SMEEB, (Municipal Ecological Environment Bureau of Shenzhen), 2018. Shenzhen
568 Pingshan River water body compliance plan.
569 http://meeb.sz.gov.cn/ydmh/zwgk/hjxw/content/post_2090474.html

570 Song, J., Li, W., 2019. Linkage between the environment and individual resilience to

- 571 urban flooding: A case study of shenzhen, china. *Int. J. Environ. Res. Public Health*
572 16. <https://doi.org/10.3390/ijerph16142559>
- 573 SRTM, (Shuttle Radar Topography Mission), 2020. The DEM data source.
574 <https://srtm.csi.cgiar.org>
- 575 Sudheer, K.P., Nayak, P.C., Ramasastri, K.S., 2003. Improving peak flow estimates in
576 artificial neural network river flow models. *Hydrol. Process.* 17, 677–686.
577 <https://doi.org/10.1002/hyp.5103>
- 578 Sudriani, Y., Ridwansyah, I., A Rustini, H., 2019. Long short term memory (LSTM)
579 recurrent neural network (RNN) for discharge level prediction and forecast in
580 Cimandiri river, Indonesia. *IOP Conf. Ser. Earth Environ. Sci.* 299.
581 <https://doi.org/10.1088/1755-1315/299/1/012037>
- 582 Sun, W., Wang, Y., Wang, G., Cui, X., Yu, J., Zuo, D., Xu, Z., 2017. Physically based
583 distributed hydrological model calibration based on a short period of streamflow
584 data: Case studies in four Chinese basins. *Hydrol. Earth Syst. Sci.* 21, 251–265.
585 <https://doi.org/10.5194/hess-21-251-2017>
- 586 SZN, (Shenzhen news), 2020. A heavy rain caused water logging in Baoan.
587 <https://mp.weixin.qq.com/s/AegwOMY7HTl66pCgZ9RDkw>
- 588 Taylor, K.E., 2001. in a Single Diagram. *J. Geophys. Res.* 106, 7183–7192.
589 <https://doi.org/10.1029/2000JD900719>
- 590 Tikhamarine, Y., Souag-Gamane, D., Ahmed, A.N., Sammen, S.S., Kisi, O., Huang,
591 Y.F., El-Shafie, A., 2020. Rainfall-runoff modelling using improved machine
592 learning methods: Harris hawks optimizer vs. particle swarm optimization. *J.*
593 *Hydrol.* 589, 125133. <https://doi.org/10.1016/j.jhydrol.2020.125133>
- 594 Vetrivel, N., Elangovan, K., 2017. Application of ANN and ANFIS model on monthly
595 groundwater level fluctuation in lower Bhavani river basin. *Indian J. Geo-Marine*
596 *Sci.* 46, 2114–2121.
- 597 Wang, S., Yao, X., 2013. Using class imbalance learning for software defect prediction.
598 *IEEE Trans. Reliab.* 62, 434–443.
- 599 Xiong, Y., Su, Z., Zhang, Y., Liang, H., 2010. Evaluation of Water Pollution of the
600 Pingshan River in Shenzhen. *Environ. Sci. Surv.* 29, 79–81.
- 601 Xu, D., Ouyang, Z., Wu, T., Han, B., 2020. Dynamic trends of urban flooding mitigation
602 services in Shenzhen, China. *Sustain.* 12, 1–11. <https://doi.org/10.3390/su12114799>
- 603 Yan, H., He, X., Lei, Y., Wang, Y., Su, H., Jiang, S., 2019. Land use-induced change in
604 trophic state of Shenzhen Bay (South China) over the past half-century. *Mar. Pollut.*
605 *Bull.* 145, 208–213. <https://doi.org/10.1016/j.marpolbul.2019.05.046>
- 606 Yang, S., Yang, D., Chen, J., Santisirisomboon, J., Lu, W., Zhao, B., 2020. A physical
607 process and machine learning combined hydrological model for daily streamflow
608 simulations of large watersheds with limited observation data. *J. Hydrol.* 590,
609 125206. <https://doi.org/10.1016/j.jhydrol.2020.125206>

- 610 Yang, S., Yang, D., Chen, J., Zhao, B., 2019. Real-time reservoir operation using
611 recurrent neural networks and inflow forecast from a distributed hydrological model.
612 J. Hydrol. 579, 124229. <https://doi.org/10.1016/j.jhydrol.2019.124229>
- 613 Yoon, H., Jun, S.C., Hyun, Y., Bae, G.O., Lee, K.K., 2011. A comparative study of
614 artificial neural networks and support vector machines for predicting groundwater
615 levels in a coastal aquifer. J. Hydrol. 396, 128–138.
616 <https://doi.org/10.1016/j.jhydrol.2010.11.002>
- 617 Zhang, D., Martinez, N., Lindholm, G., Ratnaweera, H., 2018. Manage Sewer In-Line
618 Storage Control Using Hydraulic Model and Recurrent Neural Network. Water
619 Resour. Manag. 32, 2079–2098. <https://doi.org/10.1007/s11269-018-1919-3>
- 620 Zhang, J., Zhang, Y., Song, J., Cheng, L., Kumar Paul, P., Gan, R., Shi, X., Luo, Z.,
621 Zhao, P., 2020. Large-scale baseflow index prediction using hydrological modelling,
622 linear and multilevel regression approaches. J. Hydrol. 585, 124780.
623 <https://doi.org/10.1016/j.jhydrol.2020.124780>
- 624 Zhang, W., Li, A., Jiang, X., 2013. Preliminary study on computing the area of mountain
625 regions in China based on DEM (in Chinese). Geogr. Geo-information Sci. 5, 58–
626 63.
- 627 Ziegler, A.D., Lim, H.S., Jachowski, N.R., Wasson, R.J., 2012. Reduce urban flood
628 vulnerability. Nature 481, 145.
- 629
- 630