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Egyptian sculptures from Imperial Rome. Non-destructive characterization of granitoid statues through macroscopic methodologies and in situ XRF analysis

Sander Müskens¹  · Dennis Braekmans^{2,3} · Miguel John Versluys¹ · Patrick Degryse⁴

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Abstract *Aegyptiaca*-like Domitian's obelisk is now decorating Bernini's fountain on Piazza Navona or the Egyptian lions flanking Michelangelo's stairs towards the Capitol figure prominently amidst Rome's cultural heritage. Motivations for the import, contextualization, and copying of these objects during the Imperial Roman period are as heavily debated as they are ill understood. Provenance determination plays an important role in these discussions in terms of a (supposed) dichotomy between *Egyptian* (real) versus *egyptianising* (copy) but has only been applied stylistically and never been tested analytically. A scientific characterization of the materials themselves is even lacking altogether, as is an investigation into the cultural and symbolic meaning of the materials used. This paper is a first attempt to address these important lacunae on the basis of an explorative study of a selected sample of Egyptian statues from Rome. The identification and provenance attribution of the materials used for these statues are often problematic due to their relatively fine-grained nature and dark color. Therefore, a full non-destructive analysis of Egyptian statues in dark-colored rocks

is presented in this study, with the stones evaluated by macroscopic examination and handheld X-ray fluorescence (XRF) analysis. The implemented methodology has allowed a distinction between greywacke and several varieties of granitoid rocks. In order to evaluate the potential for source attribution, a comparison was made between the results of our analyses and geochemical data for several granitoid rocks from Egypt. This has suggested Aswan as most likely source. The results presented here indicate that handheld XRF analysis can be used for the assessment of compositional variability in and potentially for the provenance of granitoid rocks, provided that a fine-grained area of the material can be measured on multiple locations, and if these values can be assessed on (in)consistencies with other published reference materials.

Keywords *Aegyptiaca* · Imperial Rome · Macroscopic rock classification · Non-destructive handheld XRF spectrometry · Provenance analysis

Introduction

Egyptian and egyptianising statues from Imperial Rome (so-called *Aegyptiaca*) form an eye-catching part of the city's cultural heritage in both the actual cityscape and Rome's museums. They testify to a process of cultural transference whereby Rome shows imperial conquest and world domination through Egyptian objects as trophies while simultaneously these (same) Egyptian objects constitute Rome as the cosmopolis by helping to build Rome's society, culture, and religion. What once was Egyptian, therefore, already soon seems to have become Roman. Besides the import of statues from Egypt, sometimes already centuries old, new sculptures with Egyptian themes were produced in the Roman world. Scholarship has traditionally understood these coexisting

✉ Sander Müskens
s.w.g.muskens@arch.leidenuniv.nl

¹ Classical & Mediterranean Archaeology, Faculty of Archaeology, Leiden University, Einsteinweg 2, 2333 CC Leiden, The Netherlands

² Materials in Art and Archaeology, Laboratory of Materials Science, Delft University of Technology, Mekelweg 2, 2628 CD Delft, The Netherlands

³ Laboratory for Ceramic Studies, Faculty of Archaeology, Leiden University, Einsteinweg 2, 2333 CC Leiden, The Netherlands

⁴ Centre for Archaeological Sciences, Department of Earth and Environmental Sciences, KU Leuven, Celestijnenlaan 200E-bus 2408, 3000 Leuven, Belgium

aspects of Egyptian sculpture in the Roman world as two essentially different phenomena. Thus, authentic *Egyptian* objects would mainly testify to Roman preoccupations with Egyptian religion and the cult of Isis in particular, while their derivative non-Egyptian and therefore less authentic counterparts, *egyptianising* copies, could also attest more generally to a Roman predilection for things exotic (Bosticco 1952; Quack 2003; Malaise 2005). Consequently, the (supposed) provenance of *Aegyptiaca* is often applied as a heuristic device to determine their archeological interpretation (for the category of *Aegyptiaca*, see Müskens 2014a). Provenance determination has, however, only been applied stylistically, based on an alleged direct relationship between cultural styles and geographic origin. The provenance of the materials themselves has not been involved in this discussion to date, despite its potential to add to the long-standing Egyptian versus egyptianising dichotomy. In fact, research on *Aegyptiaca* has so far empathically neglected the material aspects of *Aegyptiaca* in terms of both a scientific characterization of the material itself and the cultural-historical reasons for the use of particular materials.

Recent studies have shown the great potential of material culture studies for a better understanding of the socio-cultural role and impact of material culture (Degryse and Shortland 2013; Jones and Boivin 2010; Hollenback and Schiffer 2010; Brysbaert 2007). It has been demonstrated, for instance, that certain materials were sometimes deliberately used to evoke specific cultural and symbolic connotations. In the Roman world, this was particularly true for the wide range of exotically colored or patterned stones that ranked among the most sought after commodities of the Empire exactly because of the social implications of their materiality and, consequently, their potential to create specific meanings by actively capitalizing upon these implications. Many *Aegyptiaca* that circulated through the Roman world are made out of stone, and recent studies have just begun to show the relevance of a material approach for a more complete understanding of these objects (Müskens 2014b and 2017; Versluys et al. 2014; Bülow Clausen 2014). They demonstrate the necessity for a more integrated approach to *Aegyptiaca* from the Roman world. It has become clear that stylistic and iconographic analysis alone cannot provide full answers to questions about the motivations for their import, contextualization, and copying—all of which remain heavily debated and ill understood.

In order to enable a material perspective and to start exploring new directions of research, we are in need, first of all, of reliable characterizations of the materials themselves. The traditional focus on representative aspects of *Aegyptiaca* mentioned previously means that the stone materials have never been the subject of a proper analysis. As a result, there are many misidentifications in the existing literature and often geologically incorrect rock names are used in overviews like Malaise 1972, Roulet 1972, Lembke 1994, and Versluys

2002. A survey of relevant studies shows that this confusion relates in particular to more or less homogeneous, dark-colored stones. The dark stone of a male torso which is currently preserved in Palazzo Altemps in Rome is a good case in point (PA362624, Fig. 1i). It has previously been identified as “dunkles Hartgestein,” “basalto nero,” “basanite,” and, most recently, “granodiorite” (Lembke 1994; Arslan 1997; 390 V. 7 [L. Sist]; Walker and Higgs 2001, 328–329 no. 347 [C. Alfano]; Candilio et al. 2011, 324 [L. Sist Russo], respectively). The confusion between dark-colored rock types such as basalt, greywacke, and granodiorite has been widely acknowledged in Egyptian archeological literature and resonates in more general terms with the problem of incorrect characterizations of archeological stone by non-specialist archeologists (Brown and Harrell 1998; Aston et al. 2000; Klemm and Klemm 2001; Bloxam et al. 2014; on the issue in general, Herz and Garrison 1998). The Rosetta Stone is one of the most illustrative examples of this practice. Although for many years it was assumed to be made of basalt, recent analysis determined that it was actually carved from granodiorite (Middleton and Klemm 2003). Cleaning revealed that the confusion was most likely due to a protective coating and accumulated dirt which had obscured the true appearance of the rock for years. This example is illustrative for the difficulties that may be encountered in identifying archeological stone materials, which is often further complicated by unfavorable lighting conditions in museum settings. Additionally, the



Fig. 1 a–q Overview of the statues included in this study. **a** MC35. **b** TD590. **c** TD56356. **d** TD no inv. **e** MC28. **f** MC30. **g** PA362624. **h** PA362622. **i** PA362623. **j** PA60921. **k** MC31. **l** PD514563. **m** MC26. **n** MC32. **o** MC2384. **p** PA182594. **q** PA112108. Further details in Table 1

typically polished surfaces of archeological artifacts pose serious limitations to the possibilities for mineral and rock identification, especially in combination with fine-grained textures and dark colors.

Although several optical and chemical analytical methods are available to provide characterizations of and source discrimination between archeological stone materials, their specific sampling requirements often violate the nature of archeological artifacts (Kempe and Harvey 1983; Tykot 2004). This also applies to the *Aegyptiaca* in this study which require full non-destructive and in situ analysis. Therefore, we have explored macroscopic classification as described by Brown and Harrell (1991) as heuristic tool in this study. The preliminary data thus obtained were evaluated with handheld X-ray fluorescence (HH-XRF) analysis to assess the chemical variability and determine potential source areas for the materials under study. In the last decades, the development of HH-XRF devices has allowed the non-destructive and in situ determination of the chemical composition of various archeological artifacts (Shugar and Mass 2012). Many studies have looked at obsidian (Glascock et al. 1999; Frahm 2014) and other types of rocks (Barbera et al. 2013; Palumbo et al. 2015), glass (Scott et al. 2012; Scott et al. 2014), ceramics (Goren et al. 2011; Barone et al. 2011; Speakman et al. 2011; Hunt and Speakman 2015), metals (Fernandes et al. 2013), and sediments (Neff et al. 2012). This type of analysis holds great potential for the characterization of all non-moveable museum artifacts, but the results need to be carefully examined and contextualized to obtain meaningful results. In the remainder of this paper, we will explore the possibilities for full non-destructive and in situ analysis of the stone materials of a selected sample of seventeen *Aegyptiaca* from Imperial Rome. The following issues will be addressed: (1) rock classification of unknown dark-colored Egyptian statues from Rome and the potential of careful macroscopic examination with non-destructive in situ chemical analysis, (2) assessment of the validity and ability of HH-XRF to detect consistent and meaningful differences in granitoid composition, and (3) assessment of the possibility to determine an Egyptian origin for the studied rocks.

Materials: the statues

The selection of statues was primarily determined by an existing uncertainty over the identification of dark-colored rock types and the consequent need for reliable classifications of these materials in particular. Therefore, the studied sample includes seventeen *Aegyptiaca* from unknown dark-colored stone materials (Table 1 and Fig. 1a–q). The selected statues have all been found in Rome. In some cases, the Imperial Roman-use contexts are known, and it is evident that several statues once adorned the Iseum Campense, the sanctuary

dedicated to the goddess Isis on the Campus Martius (Lembke 1994). Hieroglyphic inscriptions, typology, and stylistic features suggest that the majority of the selected *Aegyptiaca* were manufactured prior to the Roman period and subsequently transported from Egypt to Rome in the Roman Imperial period. Possible exceptions are the royal male statue (PA60921, Fig. 1j) and the statue of the god Apis (PA182594, Fig. 1p) which have been variably dated to the Ptolemaic and Roman periods (La Rocca and Parisi Presicce 2010; Candilio 2011; Manera and Mazza 2001).

Analytical methods

Macroscopic rock classification and provenance hypotheses

Provisional rock classifications were formulated on the basis of the recommendations for macroscopic rock classification by Brown and Harrell (1991). Adapted from internationally acknowledged non-macroscopic analytical methods, this classification is particularly suitable for the selected *Aegyptiaca* since it meets the requirements to study these objects non-destructively and in situ. In addition, a neodymium magnet was used to test the magnetic properties of minerals in the studied rocks. This is an easy way to determine the presence of certain iron-rich minerals, most notably magnetite, which is an important asset in identifying the genetic origin of rocks (Bourne 1993). This is of particular relevance for the present study, because the magnetic susceptibility of the studied rocks can be used as a diagnostic tool to distinguish between the most frequently mistaken rock types, namely greywacke, basalt, and granodiorite. Although a wide overlap has been reported between different rock types, sedimentary rocks have the lowest average magnetic susceptibility values and basic igneous rocks have the highest. This means that greywacke, a slightly metamorphosed sedimentary rock, will be much less susceptible to the neodymium magnet than granodiorite and especially basalt and intermediate and basic igneous rocks, respectively. Telford et al. (1990) report average magnetic susceptibility values of 70 for basalt and 0.4/0.9 for sandstone/average sedimentary rocks, respectively ($\times 10^3$, SI units); and Hernant (2003) reports maximum volume susceptibility values (SI units) of 0.18 for basalt, 0.062 for granodiorite, and 0.0012/0.0209 for silt/sandstone, respectively (cf. Clark and Emerson 1991; Hunt et al. 1995).

In this paper, we use the following size scale: fine, less than 1 mm; medium, 1–5 mm; coarse, 5–30 mm; and very coarse, more than 30 mm. The terms aphanitic and phaneritic are sometimes used to determine the degree of coarseness of rocks. Aphanitic rocks are rocks in which individual crystals are not distinguishable by the unaided eye. In phaneritic rocks, crystals are visible with the naked eye. Following the

Table 1 Description of the seventeen studied sculptures

Sculpture	Proposed dating	Classification	Color index (CI)	Description	Macrogroup	Chemical group
Recumbent sphinx	MC35 Late Period (26th Dynasty); reign of Amasis, 570–526 BC	Greywacke	n/a	n/a	n/a	n/a
Stature (fragment)	TD590 No data	Greywacke	n/a	n/a	n/a	n/a
Stature of Bes	TD563.56 No data	Greywacke	n/a	n/a	n/a	n/a
Naophoros (fragment)	TD (no inv.) Late Period	Granodiorite	n.d.	Fine grained, aphanitic; non-porphyritic; pink granitic vein	1	A
Recumbent lion	MC28 Late Period (30th Dynasty)–early Ptolemaic Period	Granodiorite	n.d.	Medium to mainly fine grained, largely aphanitic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 5 mm); pink granitic veining	1	A
Recumbent lion	MC30 Late Period (30th Dynasty)–early Ptolemaic Period	Granodiorite	n.d.	Medium to mainly fine grained, largely aphanitic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 5 mm); pink granitic veining	1	A
Recumbent lion (fragment)	PA362624 Late Period (30th Dynasty)–early Ptolemaic Period	Granodiorite	CI ≈ 25% (Bt)	Medium to mainly fine grained, largely aphanitic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 5 mm); pink granitic veining	1	A
Recumbent sphinx	PA362622 Early Ptolemaic Period	Granodiorite	CI ≈ 20–25% (Hbl + Bt)	Medium to mainly fine grained, largely aphanitic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 5 mm); pink granitic vein	1	A
Male torso	PA362623 Ptolemaic Period	Granodiorite	CI ≈ 20–25% (Bt)	Medium to mainly fine grained, aphanitic; slightly porphyritic (feldspar phenocrysts up to ca. 5 mm); pink granitic vein	1	A
Royal statue	PA60921 Ptolemaic Period–Roman Period	Granodiorite	CI ≈ 20–25% (Bt)	Fine to medium grained, aphanitic/phaneritic; non-porphyritic (rare feldspar phenocrysts up to ca. 8 mm); pink granitic veining	1	A
Stature of falcon	MC31 Late Period (30th Dynasty)–Ptolemaic Period	Granodiorite	CI ≈ 20–25% (Hbl + Bt)	Fine to medium grained, aphanitic/phaneritic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 10 mm)	1	A
Lion/sphinx (fragment)	PD514563 No data	Granodiorite	CI ≈ 20–25% (Hbl + Bt)	Medium grained, phaneritic; slightly porphyritic (occasional feldspar phenocrysts up to ca. 10 mm)	1	Variant
Stature of Thoht	MC26 Late Period (30th Dynasty); reign of Nectanebo II, 360–342 BC	Granite	CI ≈ 15% (Bt)	Medium to mainly fine grained, phaneritic; non-porphyritic, gneissoid	2	B
Stature of Thoht	MC32 Late Period (30th Dynasty); reign of Nectanebo II, 360–342 BC	Granite	CI ≈ 15% (Bt)	Medium to mainly fine grained, phaneritic; non-porphyritic, gneissoid; biotite schlieren	2	B
Naophoros (fragment)	MC2384 Late Period–early Ptolemaic Period	Granite	CI ≈ 15% (Bt)	Medium to mainly fine grained, phaneritic; non-porphyritic, gneissoid	2	B
Stature of Apis	PA182594 Ptolemaic Period–Roman Period	Granodiorite	CI ≈ 20–25% (Hbl + Bt)	Coarse to mainly medium grained, phaneritic; porphyritic (feldspar phenocrysts up to ca. 30 mm, frequently exhibiting rapaviki texture); pink granitic vein	Variant 3	Variant
Head of priest	PA112108 Late Period (30th Dynasty)–early Ptolemaic Period	Granodiorite	CI ≈ 20% (Hbl + Bt)	Medium grained, phaneritic; porphyritic (feldspar phenocrysts up to ca. 25 mm), gneissoid texture	Variant 4	Variant

MC Museo Capitolino, PA Palazzo Altamps, PD Palatine depository, TD Terme di Diocleziano, n/a not applicable, n.d. not determined

recommendations in Brown and Harrell (1991), the boundary between aphanitic and phaneritic rocks is set at 1 mm which means that all fine-grained rocks are considered aphanitic. Some rocks have grains in two different size ranges. These rocks are named porphyritic, with the larger crystals called phenocrysts. The terms euhedral, subhedral, and anhedral are used to describe the degree to which crystals have developed their typical crystal morphology. In descending order, these terms indicate how well crystals are shaped, which may help in mineral identification. Alkali feldspar phenocrysts sometimes cross over into plagioclase at their rims. Macroscopically, this appears as a white mantle around a pinkish core; occasionally, plagioclase phenocrysts also cross over into alkali feldspar at their rims, which appears at a macroscopic level as a pink mantle enveloping a plagioclase crystal. This is called rapaviki texture. Igneous rocks sometimes exhibit a (sub-)parallel arrangement of the feldspar and biotite grains. This type of foliation is caused by magmatic flowage rather than metamorphism. Igneous rocks with such textures are described as gneissoid rocks. Some igneous rocks contain irregular patches or streaks which appear as portions richer in biotite than the surrounding mass and therefore darker in color or as patches of coarser or finer grains than the main rock; these are known as schlieren. Color index, that is the ratio of dark-colored minerals to light-colored minerals in a rock (Le Maitre et al. 2002), was determined by visual approximation. Color index is a useful indicator of the presence of certain types of minerals in igneous rocks and therefore an important macroscopic asset in determining the specific rock type. Color

descriptions were made according to the *Munsell Rock Color Book* (rev. ed. 2009). Where possible, potential source attributions were formulated through comparisons between the studied materials and the hand specimens of geological rock samples in the Ancient Egyptian Stone Collection (University of Toledo, Ohio; polished slabs of hand specimens from the Ancient Egyptian Stone Collection have been published online at [http://www.eeescience.utoledo.edu/faculty/harrell/Egypt/Quarries/Hardst Quar.html](http://www.eeescience.utoledo.edu/faculty/harrell/Egypt/Quarries/Hardst%20Quar.html) and will be referred to henceforth as *AESC*, followed by the numbering system used on this website) and the Klemm Collection (British Museum, London).

X-ray fluorescence analysis (HH-XRF)

Handheld X-ray fluorescence equipment (Bruker Tracer III-SD) was used to determine the chemical composition of the rocks of the selected *Aegyptiaca*. The instrument is equipped with an Rh anode X-ray tube and a Peltier-cooled silicon drift detector (~145 eV at Mn Ka). Spot size is approximately 2 by 3 mm. Because of the spot size of a HH-XRF device and homogeneity considerations, care was taken to concentrate analysis on the most fine-grained part of the different statues in order to achieve the most consistent bulk chemical data.

Measurements were taken in air for 300 s, using a Cu-Ti-Al filter, with beam conditions of 40 keV and 10.5 μ A for optimal excitation of elements from 17 to 40 keV (Fig. 2). Light elements were measured under vacuum, without a filter, and beam conditions of 15 keV and 25 μ A. An empirical

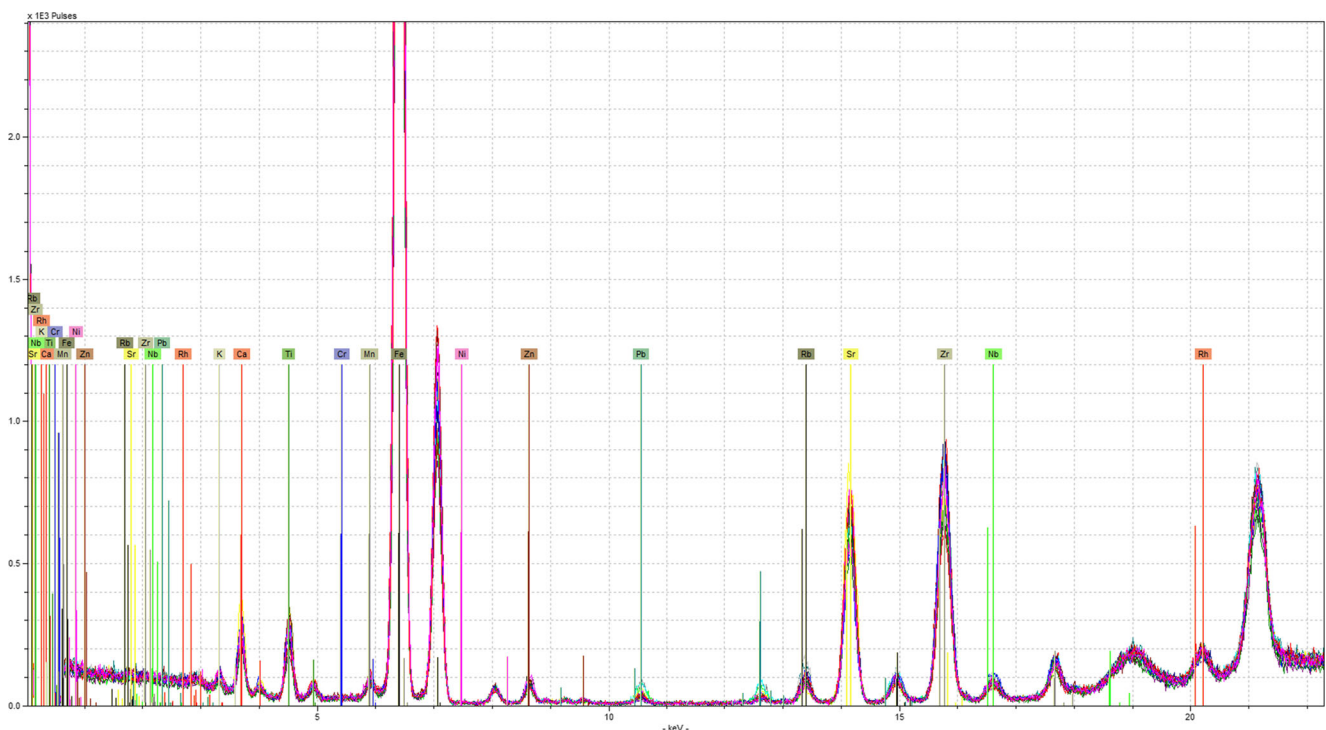


Fig. 2 Representative XRF spectrum of granodiorite group 1, measured for 300 s (40 kV–10.5 μ A) in a dry air environment

calibration was used to semi-quantitatively determine the composition of the samples. In order to check accuracy and monitor for any machine drift during the analyses, a series of rock and soil standards were also analyzed. Prior to quantifying the spectra, all data was evaluated through the ARTAX software in order to determine the consistency of the matrices. A set of international certified standards was used to determine accuracy: BIR-1 (basalt), SRG-1 (shale), GSP-2 (granodiorite), 2710a (soil), 98b (sediment), and CRM667 (sediment). Only elements with sufficiently high squared correlation coefficients (R^2) (intensities/certified value), as an assessment of accuracy, were retained for subsequent analysis: Ca = 0.90, Sr = 0.96, Ti = 0.99, Mn = 0.99, Fe = 0.98, Ni = 0.91, Zn = 0.99, Zr = 0.98, Cr = 0.93, and K = 0.97. Other elements did not provide any acceptable coefficients and were therefore not taken into account for the analyses. Precision (both repeatability and reproducibility) of the measurements was controlled at several instances by replicate analyses (no. 5) and is best assessed through the calculation of the relative standard deviation (RSD or %RSD) (Abzalov 2008). All elements are well below 10% RSD: Ca (1.78), Fe (0.88), K (4.87), Sr (1.33), Ti (3.41), Zn (5.78), Cr (3.19), and Zr (1.53), apart from Ni (9.97) (GSP2 and BIR1a).

The measurements were evaluated by an assessment of semi-quantitative data through bivariate diagrams as well as by means of multivariate statistical procedures such as principal component analysis (henceforward PCA). These statistical techniques were selected in order to structure the data and to explore potential chemical factors contributing to the variability between the statues (Davis 1986). All statistical procedures were carried out with the Statistica software (version 8.0).

Results and discussion

Macroscopic rock classification and provenance hypotheses

The rocks of statues MC35, TD590, and TD56356 were found to be essentially different from all others in the studied sample. They are fine-grained, aphanitic rocks with very dense, homogeneous matrices. MC35 is olive black, and TD590 and TD56356 are dark gray. Due to their fine-grained nature, exact grain sizes and mineralogy could not be determined. No visible attraction between the neodymium magnet and these rocks could be observed. This and the other macroscopic characteristics are indicative of greywacke from the Wadi Hammamat in Egypt, the only known ancient quarry for this rock type (Bloxam et al. 2014; Brown and Harrell 1995). The rocks from this location are slightly metamorphosed, compact sedimentary rocks with abundant clay/mica that texturally varies from sandstone (predominant grain size 0.062–2 mm) to mudrock (0.004–0.062 mm). Their colors range from dark

gray to nearly black and greenish gray to grayish green (cf. AESC 28a (a) variety 2 and AESC 28a (a) variety 1, respectively). Pale yellowish brown rounded clasts are visible on the right flank of MC35 (diam. ca. 10 and 3 cm, respectively, i.e., falling within the cobble and pebble size range). Comparable clasts can be observed on several artifacts carved from the Wadi Hammamat greywacke (De Nuccio and Ungaro 2002, 341 no. 41 [P. Liverani]; De Caro 2006, 202 no. III.108 [R. Pirelli]). Based on the strong macroscopic analogies with greywacke from the Wadi Hammamat, the raw materials of statues MC35, TD590, and TD56356 are likely to originate from this Eastern Desert source.

Igneous plutonic origins were determined for all of the remaining fourteen rocks in the sample. Granularity could be observed by the unaided eye in most cases (i.e., these are phaneritic rocks), which means that the average grain size is above ~1 mm. The majority of the statues have well developed textures that are indicative of their plutonic origin. In most cases, the attraction between the studied rocks and the neodymium magnet could be clearly observed. The color indexes, as far as these could be established by visual approximation, range between ~15–25%, and the overall rock colors vary between different shades of gray. This is an indication for the felsic to intermediate compositions of these fourteen rocks. More specifically, the relative abundance of quartz and alkali feldspar relative to biotite and hornblende suggests that the studied granitoid rocks compositionally range from granite to granodiorite. There are, however, several textural and compositional variations among the fourteen statues. Based on this variability, two macroscopic groups with similar appearing stones were recognized, group 1 with nine statues and group 2 with three statues, and another two statues are carved from stones that are dissimilar to all other stones in this study. The latter three stones are referred to as variants 3 and 4 (see Table 1).

Group 1 is the largest group with nine statues. The rocks in this group are dark gray and typically appear as grayish black in polished surfaces. Color indexes are approximately 20–25%. These rocks are fine to medium grained and have overall fairly homogeneous granular matrices. The finer-grained specimens are largely aphanitic, although some grains can be distinguished with the naked eye, especially on broken surfaces and at a suitable angle to catch the light on cleavage faces. These rocks are therefore medium to mainly fine grained (TD no inv., MC28, MC30, PA362624, PA362622, PA362623). Feldspar phenocrysts are occasionally present and reach up to ca. 5 mm in the finer-grained specimens and ca. 10 mm in case of the fine- to mainly medium-grained statues (PA60921, MC31, and PD514563). The dark-colored matrices of seven rocks in this group are crosscut by coarse- to mainly medium-grained, very pale orange to grayish orange pink veins of granitic composition (quartz and alkali feldspar; Fig. 3a).

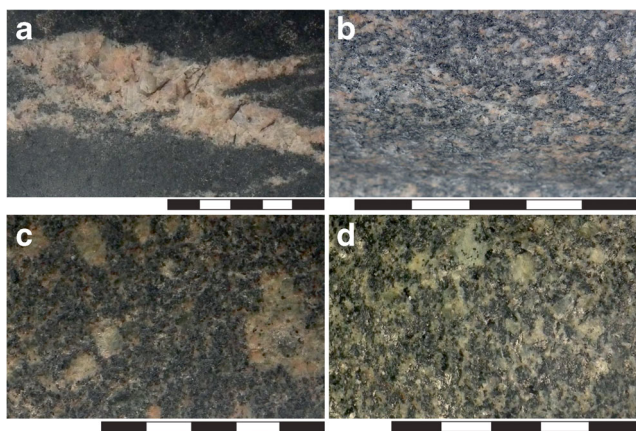


Fig. 3 a–d Macrophotographs of typical facies of studied granitoid rocks. **a** Group 1. **b** Group 2. **c** Variant 3. **d** Variant 4. Scale in centimeters

The three rocks in group 2 have a lower overall color index ($CI \approx 15\%$) and rock color. These rocks are mainly medium gray to medium light gray, but they grade in parts into medium dark gray to light gray on account of local variations in the concentrations of biotite. They are medium- to mostly fine-grained rocks with fairly equigranular textures, and they show foliation, as evidenced by the parallel arrangement of the feldspar and biotite flakes. These rocks, in other words, have a gneissoid texture (Fig. 3b). Dark-colored patches appear as a streak on front of the base of statue MC32 and as a wavy band on the right shoulder of the baboon. These biotite schlieren, which follow the direction of foliation, are richer in dark-colored biotite than the surrounding rock which accounts for their darker (dark gray to grayish black) color. As opposed to the rocks of group 1 and the two variants described in the following sections, the rocks of group 2 only weakly reacted to the proximity of the neodymium magnet.

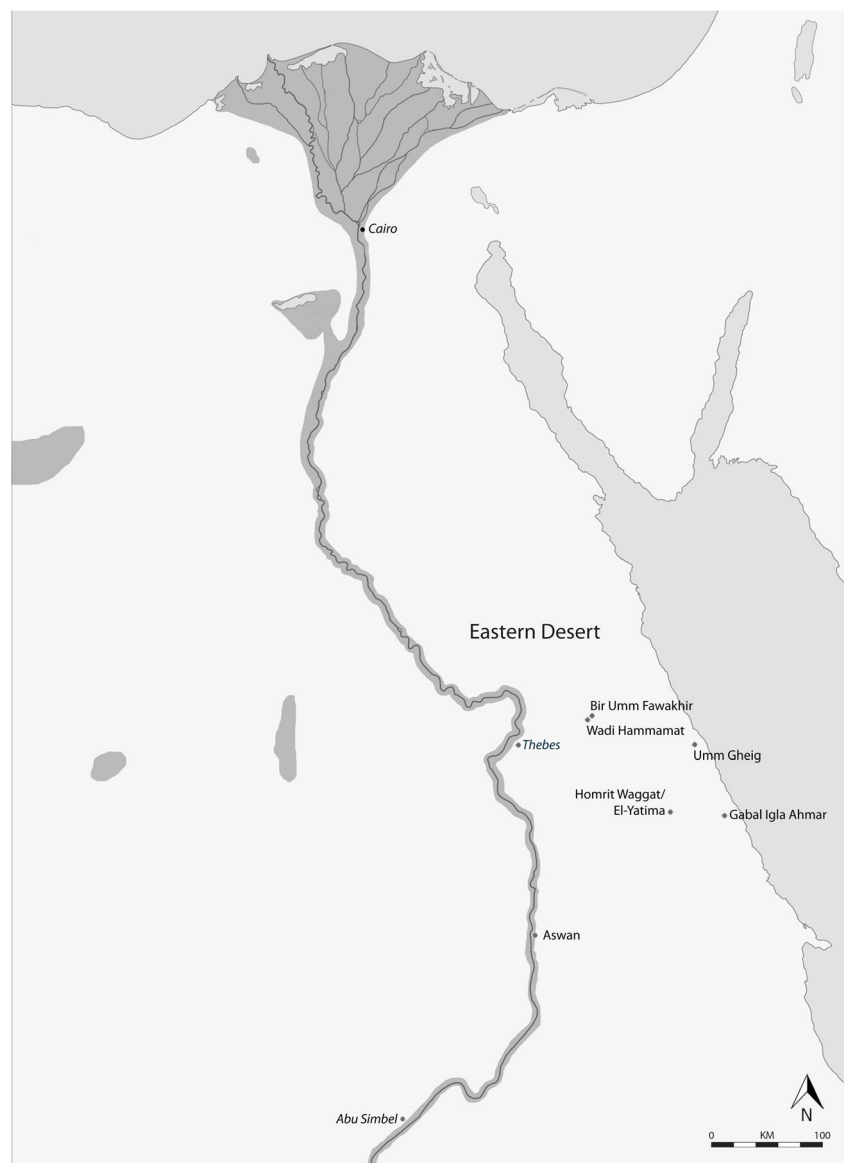
The macroscopic characteristics of the rock of the Apis statue (PA182594) are markedly different from the others in the studied sample, and, therefore, this statue is designated to variant 3 (Fig. 3c). The overall rock color is grayish black, and the color index is approximately 20–25% (hornblende and biotite can be easily observed due to large grain size). It is a coarse- to mainly medium-grained porphyritic rock with abundant anhedral to subhedral plagioclase feldspar phenocrysts up to ca. 30 mm across, and less frequent alkali feldspar phenocrysts (up to ca. 15 mm across), several of which exhibit a rapaviki texture. A medium-grained granitic vein cuts across the dark-colored matrix. The rock of the head of a priest (PA112108), finally, is another variety, variant 4 (Fig. 3d). It is an overall mottled dark gray and yellowish gray, mainly medium-grained porphyritic rock with abundant plagioclase feldspar phenocrysts up to ca. 25 mm across and $CI \approx 20\%$ (hornblende and biotite). The mostly anhedral to subhedral phenocrysts show a distinct parallel orientation. This rock is therefore a gneissoid variety of granodiorite.

A preliminary geological study has shown that strong macroscopic analogies exist between the raw materials of the fourteen statues and different granitoids outcropping in the Aswan area (Fig. 4). These rocks exhibit a wide range of compositional and textural variations, including two main varieties of granite and at least three principal types of granodiorite (El-Shazly 1954; Attia 1955; Higazy and Wasfy 1956; Aston et al. 2000; Klemm and Klemm 2008). Among these is a medium- to mainly fine-grained, non-porphyritic granite, also known as Saluja-Sehel Granite (Finger et al. 2008). These rocks vary from red/pink to gray in color, with the gray variety mainly located at the northeast of the Aswan Dam (Soliman 1980). The biotite flakes, i.e., the dominant dark-colored mineral in these rocks, often show a parallel arrangement (i.e., these are gneissoid granites). The biotite contents moreover may exhibit local variations due to which the overall rock color may vary over small distances (Attia 1955), and biotite schlieren and granitic veins are commonly observed in these rocks (Gindy 1956; Higazy and Wasfy 1956).

The most abundant variety of granodiorite at Aswan is (1) gray in color and spotted with white and pinkish feldspar phenocrysts up to ca. 30 mm across, which may be parallel aligned. A second, basic variety is (2) dark gray in color, with abundant dark-colored minerals and less well-developed feldspar phenocrysts. This includes a fine-grained variant with occasional feldspar phenocrysts up to max. ca. 2 mm across (Middleton and Klemm 2003). The third variety (3) is a gneissose granodiorite, which is often developed at Aswan near the contact with coarse-grained granite (El-Shazly 1954; Attia 1955; Noweir et al. 1990). The presence of pink granitic veins cutting across the dark-colored matrices (De Putter and Karlshausen 1992; Middleton and Klemm 2003) and the rapaviki texture of the feldspar phenocrysts (Higazy and Wasfy 1956; Ragab et al. 1978; Meneisy et al. 1979) are common features in granodiorites from Aswan.

The macroscopic characteristics of the rocks in group 1 closely correspond to the descriptions of granodiorite variety (2) from the literature. The finer-grained specimens in this group show strong similarities to *AESC* 5(b) variety 1, samples 1–2. The macroscopic features of the rocks in group 2, next, are fully consistent with published descriptions of the gray Saluja-Sehel Granite (cf. *AESC* 5 (d) variety 2, sample 1–2). Strong macroscopic analogies exist between granodiorite variant 3 and variety (1) from the literature, and variant (4) is consistent with the description of granodiorite variety (3) (cf. sample 439 in the Klemm Collection: Klemm and Klemm 2008, plate 81). Based on the strong macroscopic similarities between the studied rocks and granitoids from Aswan, it is our hypothesis that the raw materials of all fourteen statues were possibly extracted from the ancient granite-granodiorite quarries at Aswan.

Fig. 4 Map of Egypt, showing the location of sites mentioned. Names in *italics* are displayed for reference purposes



X-ray fluorescence analysis (HH-XRF)

The chemical compositions of 38 XRF measurements on the seventeen statues are reported in Table 2. Macroscopic analysis has previously suggested that three statues in the studied sample were carved from greywacke (MC35, TD590, TD56356) and the other fourteen from granitoid rocks. To evaluate these observations, a first multivariate analysis of all chemical elements by PCA was conducted to cover and identify potential geochemical variation. A graphical output shows that the first two components cover ~60% of the variability (Fig. 5). This plot shows, first of all, a clear separation of greywacke from the other studied rocks, which is mainly due to lower values of Zr, FeO (total), and TiO₂. There is, however, also significant variation detected between the different measurements of the granitoid stones. In order to

evaluate if and to what extent this variability corresponds to the potential identification of different groups of granitoid rocks and their varieties on the basis of macroscopic examinations, and to assess the provenance hypotheses formulated previously, in the remainder of the analyses, we will focus only on the compositional variability in granodiorite.

Oxide values of the granitoid measurements are reported in wt.%, all others in ppm. A brief overview of the analytical output reveals CaO lower and upper quartile ranges between 2.64 and 4.07 wt%. Only one individual measurement is above 8.0 wt% (PA182594). The total FeO content has a rather broad range, from 2.40 to 8.40 wt% lower and upper quartile. K₂O is mostly restricted within the 0.9 to 1.87 wt% range. TiO₂ is very variable between the different statues, with samples on the lower end 0.77 and ~2.0 wt% on the high end. Trace element composition is relatively homogeneous with

Table 2 HH-XRF analyses of sampled sculptures

Sculpture	Classification	K ₂ O (wt%)	CaO (wt%)	TiO (wt%)	Cr (ppm)	MnO (wt%)	FeO (T) (wt%)	Ni (ppm)	Zn (ppm)	Sr (ppm)	Zr (ppm)	Nb (ppm)	Ba (ppm)
Z		19	20	22	24	25	26	28	30	38	40	41	56
1 MC35	Greywacke	1.16	1.52	0.43	81	0.18	4.91	66	79	253	126	4	1481
2 TD590	Greywacke	0.90	2.33	0.21	n.d.	n.d.	3.94	10	110	202	87	4	1331
3 TD56356	Greywacke	1.09	0.70	0.14	<10	n.d.	2.81	10	95	176	86	5	893
4 TD (no inv.)	GD group 1	1.17	4.15	1.54	n.d.	n.d.	8.62	n.d.	124	276	300	28	1487
5 TD (no inv.)	GD group 1	0.41	6.21	0.65	<10	n.d.	6.22	n.d.	242	296	331	29	1565
6 TD (no inv.)	GD group 1	0.57	7.19	1.02	57	n.d.	8.03	n.d.	156	273	286	26	1193
7 MC28	GD group 1	1.90	2.69	2.37	n.d.	0.13	10.69	n.d.	115	299	319	25	1465
8 MC28	GD group 1	1.06	2.55	1.69	n.d.	0.19	9.10	n.d.	180	269	288	21	1227
9 MC30	GD group 1	2.43	3.24	2.06	n.d.	0.15	9.54	n.d.	169	252	267	21	971
10 MC30	GD group 1	1.60	2.56	2.06	n.d.	n.d.	8.28	<10	105	233	238	21	1058
11 PA362624	GD group 1	1.35	3.59	2.63	100	n.d.	9.82	n.d.	116	298	282	22	1214
12 PA362624	GD group 1	2.19	4.44	2.57	59	n.d.	10.41	n.d.	112	296	303	23	1194
13 PA362624	GD group 1	1.81	2.59	1.58	n.d.	n.d.	7.44	<10	107	288	305	16	1226
14 PA362624	GD group 1	1.48	4.66	2.19	11	0.14	9.02	n.d.	110	235	299	31	987
15 PA362624	GD group 1	1.17	3.81	2.08	n.d.	n.d.	9.34	17	105	317	255	22	1059
16 PA362622	GD group 1	0.41	2.97	2.08	n.d.	0.13	7.13	16	97	290	261	19	1245
17 PA362622	GD group 1	0.74	2.68	1.94	n.d.	n.d.	6.56	n.d.	84	248	275	18	948
18 PA362622	GD group 1	1.16	3.12	1.54	<10	n.d.	7.34	n.d.	97	305	289	17	1255
19 PA362622	GD group 1	1.08	3.45	2.01	n.d.	n.d.	8.19	12	99	285	285	21	1266
20 PA362622	GD group 1	0.86	2.62	1.08	n.d.	0.13	6.13	n.d.	89	245	212	17	859
21 PA362622	GD group 1	0.90	2.79	1.34	28	n.d.	7.52	n.d.	104	257	281	18	936
22 PA362623	GD group 1	0.81	3.78	1.68	21	n.d.	8.09	n.d.	108	369	458	21	1259
23 PA362623	GD group 1	0.49	2.76	1.45	n.d.	n.d.	6.39	<10	98	269	216	18	709
24 PA362623	GD group 1	0.82	3.02	1.41	n.d.	0.12	6.81	n.d.	89	317	186	16	917
25 PA60921	GD group 1	1.43	3.48	1.37	n.d.	n.d.	7.58	n.d.	114	275	295	28	1504
26 PA60921	GD group 1	0.90	3.43	1.12	n.d.	n.d.	7.65	n.d.	170	285	307	25	1390
27 MC31	GD group 1	0.93	5.88	2.59	n.d.	0.13	8.40	n.d.	104	351	227	15	933
28 PD514563	GD group 1	1.16	4.07	1.93	n.d.	n.d.	7.17	n.d.	115	356	114	16	1132
29 PD514563	GD group 1	1.18	3.13	1.45	n.d.	0.12	6.00	n.d.	108	471	124	12	1166
30 MC26	Granite group 2	2.13	4.56	0.24	n.d.	0.04	0.89	<10	67	219	230	20	2479
31 MC32	Granite group 2	2.22	4.21	0.10	42	0.05	3.48	n.d.	108	203	382	12	2082
32 MC2384	Granite group 2	2.04	1.33	0.04	n.d.	0.33	1.17	11	52	146	336	8	1579
33 PA182594	GD variant 3	2.62	2.26	0.72	30	n.d.	4.01	n.d.	104	297	190	17	1502
34 PA182594	GD variant 3	1.44	3.87	0.99	25	n.d.	6.20	<10	116	251	342	21	1344
35 PA182594	GD variant 3	1.87	3.75	0.77	n.d.	0.09	4.99	n.d.	162	323	251	18	1664
36 PA112108	GD variant 4	1.84	2.37	0.68	<10	n.d.	3.94	n.d.	92	331	55	17	1407
37 PA112108	GD variant 4	1.39	2.21	0.67	42	0.10	3.83	n.d.	93	306	143	17	1368
38 PA112108	GD variant 4	1.94	2.64	0.71	n.d.	n.d.	4.83	n.d.	105	320	61	18	1550

n.d. not determined, *GD* granodiorite

quartile ranges between 252 and 317 ppm (Sr), 216–303 ppm (Zr), 97–116 ppm (Zn), and 1058–1487 ppm (Ba). The measurements of Cr and Ni were for the most part not detected or close to the detection limit of the instrument. Given the higher error margins, the results for these two elements are only partial in nature and should not be considered reliable enough for identification purposes.

An important feature is to determine the intra-measurement variability of the various statues through multiple spot analyses, in order to assess the ability to obtain comparable results. This was carried out on flat surfaces and concentrated on the most fine-grained parts of the matrices. The actual observations are also graphically represented in a series of biplots (Fig. 6). In general, intra-statue variability is fairly restricted,

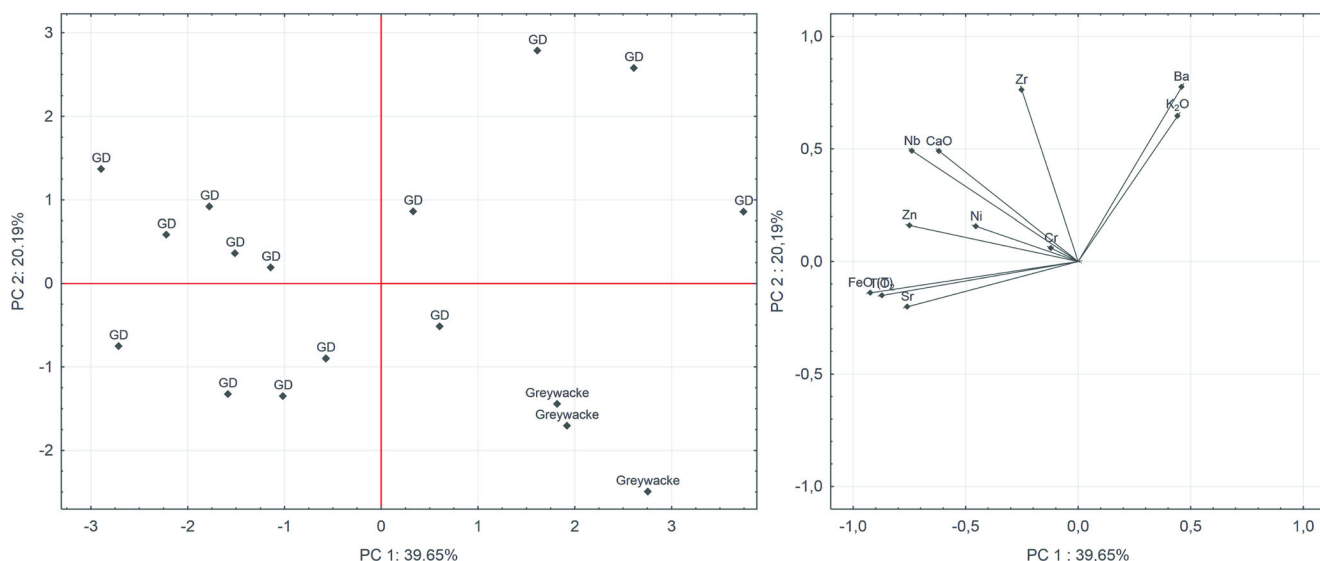


Fig. 5 Principal component analysis of seventeen statues according to rock classification. *GD* granodiorite

except for some statues where a larger range of values is found, e.g., in PA362622, values for TiO_2 range between approx. 1–2 wt%, and Sr values range in PD514563 between approx. 350–450 ppm. Although variability remains within these boundaries, a considerable overlap can be identified between most of the statues, especially related to the macroscopic grouping. This relation to the previously identified macroscopic groups is represented in Fig. 7 (see also Table 1). Macroscopic group 1 (granodiorite) comprises fine- and medium-grained material and is shown to have the widest range of chemical values. Nevertheless, these can be clearly and consistently separated from macroscopic groups 2 (granite) and variant 4 by their systematically higher Ti, Sr, and Fe values and lower Ba and K levels. Variant 3 cannot be clearly discerned chemically from macroscopic groups 1 and variant 4, yet shows considerable variation opposite to macroscopic group 2.

Principal component analysis was subsequently carried out to determine the influence and variability of all elements on the obtained average values for each statue. As such, inter-statue differences can be identified as well as the contributions

of each element towards this difference. Overall, the first three principal components cover approximately 85% of explained variance. A graphical representation of the first two components reports the construction of at least two chemical groups (Fig. 8). Most of the statues of macroscopic group 1 share a common geochemical profile of which some exhibit differences, notably statues MC31 and TD (no inv.). Statues MC2384, MC32, and MC26 form a closely connected second group and are substantially enriched in K_2O and Ba, which is consistent with the characterization of these rocks as granites. Some statues show variability opposite to macroscopic groups 1 and 2. Statue PA112108 (macroscopic variant 4), for example, shows more depleted levels of Zr in the matrix opposite to other statues. The Apis statue (PA182594, macroscopic variant 3) indicates a better chemical consistency to the main group 1. Statue PD514563, lastly, seems significantly more enriched in CaO and Sr, suggesting another potential different chemical variation.

When evaluating the elements contributing to the overall variability, a close overlap can be seen in K_2O and Ba, likely relating to a feldspar component, CaO-Sr, common for

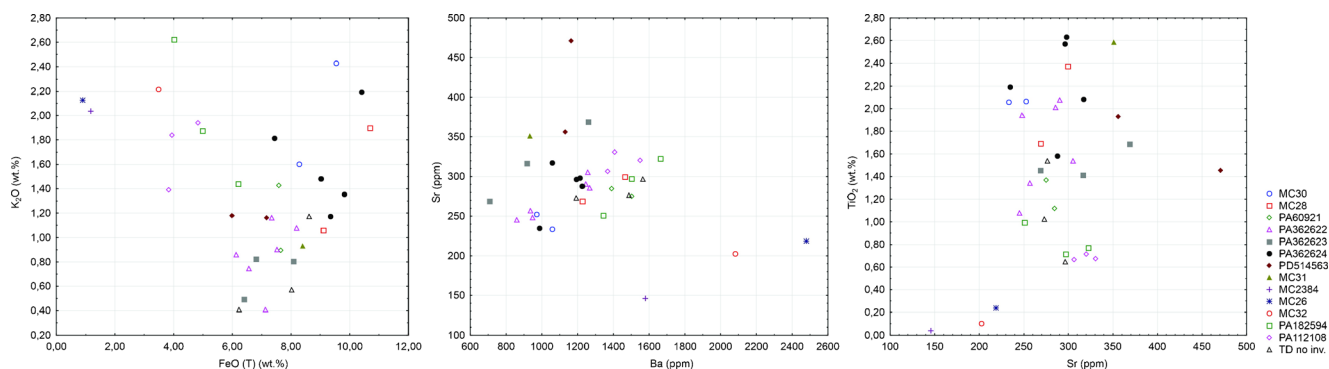


Fig. 6 Bivariate plots of individual measurements grouped per statue

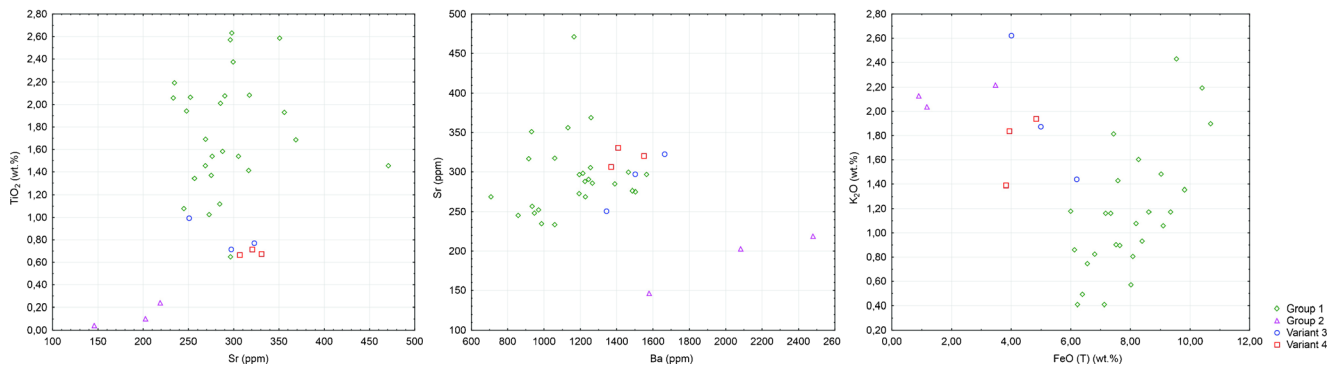


Fig. 7 Bivariate plots of individual measurements grouped per macroscopic group

carbonate phases, and FeO-TiO₂, representing a heavy mineral fraction. Other elements, such as Zr, do not seem to correlate with any other elements.

Results obtained from both the bivariate plots and statistical analyses show a consistent pattern. It can be concluded that the granodiorites (group 1) and granites (group 2) provide two consistent chemical groupings, which signify a clearly different measured chemical composition of the granitoid micro-structure. Apart from both these macrogroups, the individual samples (PA182594, PA112108, PD514563) show variability that cannot be directly attributed to either of these two groups. A clear correspondence is thus obtained between both the macroscopic and the chemical approach apart from the previously defined differentiation in chemical values for these three statues. This demonstrates the potential of discrimination on a group level when focusing on fine-grained matrices through multiple spot analysis. Based on these results, it can be concluded that HH-XRF can be a useful additional technique in exploring chemical differences for compositional studies of granitoid rocks, although with a more limited resolution of differentiation than invasive laboratory techniques like ICP-MS and NAA (Pollard et al. 2007). Laser ablation systems coupled to ICP-MS could be an alternative on a microinvasive

scale (Dussubieux et al. 2016), but still these techniques require a small sample which may be difficult to obtain, as in the case of the sculptures studied in this paper. Next, in order to assess the potential of this technique for the determination of an Aswan provenance and to evaluate macroscopic hypotheses, the chemical results will be compared to a dataset of published results from Egyptian granitoid rocks.

Discussion of provenance hypotheses

In order to evaluate the hypothesis that the raw materials of the fourteen studied granitoid statues were extracted from the granite-granodiorite quarries at Aswan, we compared the results of our analyses to previously published whole-rock analyses of granites and granodiorites from Aswan, as well as several sources of granodiorite in the Eastern Desert. Aswan granodiorites can be distinguished from stones with similar compositions from the Eastern Desert on the basis of notable differences in Fe, Ti, Ca, and K contents. Relative to granodiorites from the Eastern Desert, Aswan granodiorites are significantly enriched in total Fe and Ti contents (~8.0 wt% FeO, ~2.50 wt% TiO₂ for Aswan and ~3.5 wt% FeO, ~0.50 wt% TiO₂ for average Eastern Desert values) and relatively

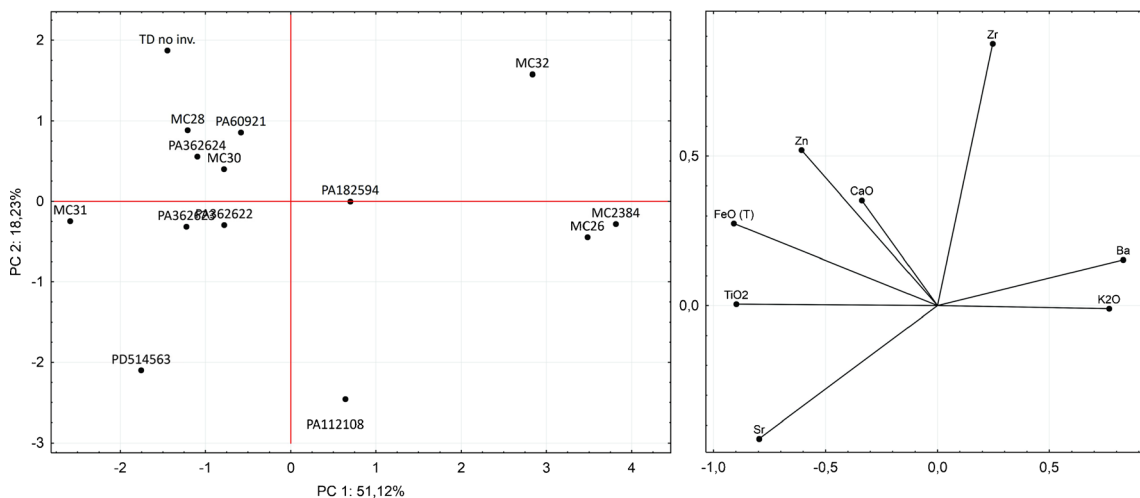


Fig. 8 Graphical distribution of extracted principal component 1 and 2 loadings in combination with the PC1 and PC2 scores

depleted in Ca and K contents. A ternary plot showing the relative Ti-Fe-K elemental composition of the sampled statues demonstrates that most of the granodiorites are relatively consistent with data from the Aswan area and different from Eastern Desert sources (Fig. 9, values are reported in Table 3; for the mentioned sites, see Fig. 4). This close chemical consistency suggests that the raw materials for these statues were indeed obtained from the quarries at Aswan.

A minority of the sculptures in this study show a different pattern. Statues MC26, MC32, and MC2384 were found to be different from the granodiorite group, which corresponds well with the macroscopic characterization of these rocks as granite. These statues are chemically characterized by notably low concentrations of FeO and TiO₂ that are not consistent with the known granodiorites from Aswan. The ternary plot in Fig. 9 shows that these rocks have a chemical composition that better matches the published data of granodiorites from the Eastern Desert and those of the Aswan fine-grained and (monzo)granites (e.g., lower TiO₂, higher K₂O) under the current analytical setup.

Of the granodiorite sources in the Eastern Desert plotted in Fig. 9, Bir Umm Fawakhir is the only one that was anciently quarried. It is, however, highly unlikely that this source provided the raw materials for the statues of group 2: MC26, MC32, and MC2384. Not only are there macroscopic differences between the rocks of these three statues and the stones from the Bir Umm Fawakhir quarries, particularly in terms of mineralogy, color, and texture, but also the quarries at Bir Umm Fawakhir were exclusively worked during the first and second centuries AD and only for the production of

architectural elements such as columns and pavement tiles (Brown and Harrell 1995; Lazzarini 2002; for a color photograph of a representative slab from Bir Umm Fawakhir, see *AESC* 27, sample 1). The fact that the relevant objects are sculptures that were carved long before the extraction at the Fawakhir quarries began eliminates these quarries as likely a source area. It is therefore highly likely that these three statues are indeed carved from granite, and not granodiorite, as the preliminary macroscopic analysis has suggested. The chemical correspondence to the fine-grained granite from Aswan, i.e., low concentrations of FeO and TiO₂ and high K₂O and CaO values, suggest that like the studied granodiorites also the fine-grained granites of macrogroup 2 were obtained from the quarries at Aswan.

Conclusions

This paper represents the first exploration of a fully non-destructive characterization of Egyptian statues from Imperial Rome. The results presented here show that careful macroscopic and chemical observations provide a clear discrimination between commonly confused stone types like greywacke, gray granite, and granodiorite. Combining a macroscopic approach with in situ HH-XRF analysis demonstrates the potential to identify granitoid rocks and their variants, although restrictions on resolution and details of analysis apply. Three main issues were addressed in this paper: (1) the relation between macroscopic and chemical analyses, (2) potential and resolution of HH-XRF for identification of

Fig. 9 Ternary diagram of samples in this study (group 1—granodiorites: *green diamonds*; group 2—fine-grained granites: *red diamonds*; variant 3—*yellow diamond*; variant 4—*black diamond*) compared to existing literature data of Aswan granitoid rocks: granodiorite (*blue squares*), tonalite (*purple squares*), monzogranite (*orange squares*), coarse granite (*black squares*), and fine granite (*green squares*). Granodiorites from the Eastern Desert and a global average composition of granodiorite are indicated in *red circles* and a *yellow square*, respectively. All values are reported in Table 3

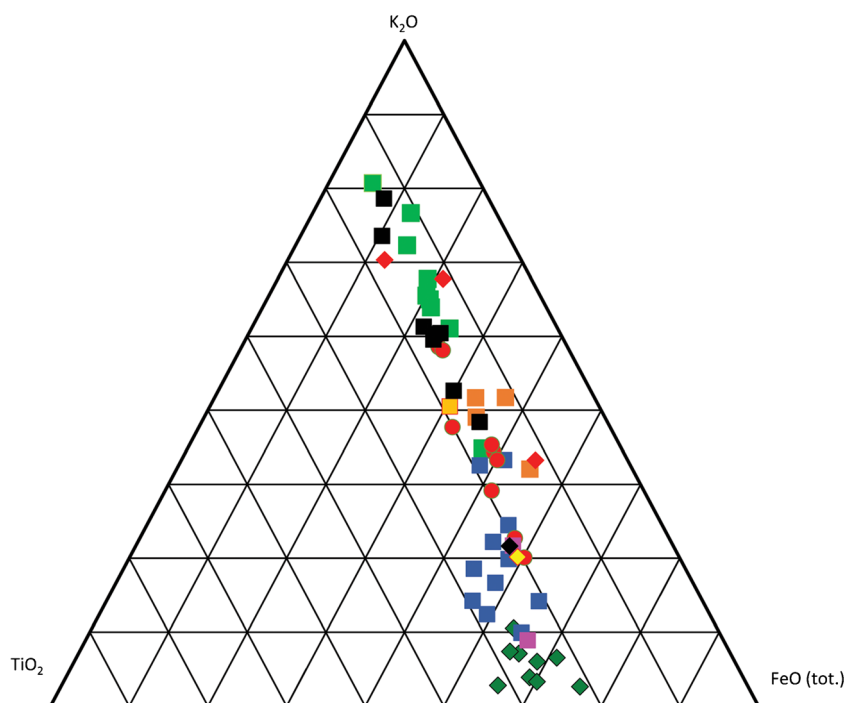


Table 3 Comparison of granitoid compositions from the present study and the literature

Sculptures (this study)	CaO (wt%)	K ₂ O (wt%)	TiO ₂ (wt%)	FeO (T) (wt%)	Macrogroup
TD no inv.	5.85	0.72	1.07	7.62	1
MC28	2.62	1.48	2.03	9.90	1
MC30	2.90	2.02	2.06	8.91	1
PA362624	3.81	1.53	2.13	8.95	1
PA362622	2.94	0.86	1.66	7.15	1
PA362623	3.20	0.83	1.65	7.51	1
PA60921	3.45	1.16	1.24	7.62	1
MC31	5.88	0.93	2.59	8.40	1
PD514563	3.60	1.17	1.69	6.58	1
MC26	4.56	2.13	0.24	0.89	2
MC32	4.21	2.22	0.10	3.48	2
MC3284	1.33	2.04	0.04	1.17	2
PA182594	3.29	1.98	0.83	5.06	Variant 3
PA112108	2.62	1.82	0.74	4.27	Variant 4
Literature values	CaO (wt%)	K ₂ O (wt%)	TiO ₂ (wt%)	FeO (T) (wt%)	
Aswan—granodiorite					
Higazy and Wasfy (1956) (<i>n</i> = 3)	4.55 ± 1.04	3.35 ± 0.75	1.70 ± 0.60	8.67 ± 2.21	—
Meneisy et al. (1979) (<i>n</i> = 5)	4.45 ± 0.67	3.07 ± 0.36	1.92 ± 0.52	8.03 ± 0.74	—
Middleton and Klemm (2003) Rosetta Stone (avg. three analyses)	4.00	2.30	1.30	8.40	—
Middleton and Klemm (2003) 489 (avg. two analyses)	4.20	2.50	2.50	8.00	—
Noweir et al. (1990) 200–201 (<i>n</i> = 2)	2.82 ± 0.29	3.57 ± 0.65	0.53 ± 0.13	5.21 ± 0.73	—
Aswan—tonalite					
Finger et al. (2008) D1	5.81	1.97	2.11	10.09	—
Finger et al. (2008) D3	4.00	3.17	1.23	7.47	—
Aswan—fine-grained granite					
Finger et al. (2008) Fg2	1.05	5.14	0.25	1.40	—
Meneisy et al. (1979) (<i>n</i> = 6)	1.56 ± 0.31	5.16 ± 0.22	0.34 ± 0.04	3.00 ± 0.45	—
Higazy and Wasfy (1956) C (grayish pink sample)	2.53	4.17	0.80	5.50	—
Higazy and Wasfy (1956) D (red sample)	1.20	7.55	0.09	2.90	—
Aswan (coarse-grained)—granite					
El-Gaby (1975) no. 35	2.13	4.23	0.66	3.99	—
Finger et al. (2008) (<i>n</i> = 2)	1.98 ± 0.28	4.83 ± 0.08	0.56 ± 0.04	3.26 ± 0.06	—
Higazy and Wasfy (1956) (<i>n</i> = 2)	1.87 ± 0.57	5.65 ± 1.06	0.43 ± 0.13	2.69 ± 0.59	—
Meneisy et al. (1979) 6–19 (<i>n</i> = 13)	1.58 ± 0.22	5.34 ± 0.59	0.50 ± 0.08	3.79 ± 0.57	—
Aswan—monzogranite					
Noweir et al. (1990) (<i>n</i> = 5)	2.57 ± 0.27	4.02 ± 0.61	0.38 ± 0.23	4.50 ± 0.77	—
Eastern Desert—granodiorite					
Bir Umm Fawakhir (Brown and Harrell 1995)	2.16	3.45	0.41	2.58 ^a	—
Gabal Igla Ahmar (Hassanen et al. 1996) (<i>n</i> = 2)	3.50 ± 0.03	2.66 ± 0.32	0.52 ± 0.03	3.35 ± 0.76	—
Homrit Waggat/El-Yatima area I (Moghazi et al. 1999) (<i>n</i> = 4)	3.81 ± 0.87	2.07 ± 0.33	0.55 ± 0.16	3.56 ± 1.05	—
Umm-Gheig (El-Sayed et al. 2002) (<i>n</i> = 2)	3.90 ± 0.13	2.8 ± 0.23	0.46 ± 0.02	3.86 ± 0.16	—
Granodiorite (global average) (Cox et al. 1979)	3.83	2.73	0.54	2.73	—

Literature values are given as mean value and standard deviation when applicable

n.d. not detected

^a Only Fe₂O₃ given

granitoid rocks, and (3) determining an Egyptian (Aswan) origin for the studied raw materials. Our conclusions are as follows:

1. Results from both methods overlap quite significantly and support the obtained grouping and separation between granite and granodiorite. While macroscopic data generally focuses on the visible macrostructure of rocks, HH-XRF is able to determine the chemical composition of a chosen part of this structure. In this paper, we explored the potential to measure the composition of the dark-colored component. Two main groups and two individual variants were discerned on the basis of macroscopic observations, while another variant was recognized on the basis of the chemical data (statue PD514563). Variants 3 and 4 show considerable differentiation in chemical features in comparison to group 2 (granites). These variants appear to be variations of granodiorite.
2. HH-XRF is a suitable non-destructive technique to explore chemical variations in granitoid rocks, bearing in mind element restrictions and ideally closely connected to macroscopic observations. The adopted approach has shown that granodiorite, granite, and greywacke can be clearly distinguished on the basis of non-destructive compositional analysis. Additionally, analysis of the studied granitoid statues has provided coefficients of variation between multiple spot measurements that are systematic at approximately ~10%. Within these margins, consistent results were obtained which allowed the relative discrimination between groups of different granitoid rocks and their variations. While these margins indicate that obtained results should be considered only as indicative when applied for direct provenancing and comparison, the results nevertheless demonstrate that granitoid materials can be discriminated using HH-XRF when fine-grained matrices are selected for analysis. This technique therefore has an explorative grouping potential that needs further consideration, in particular, because it may help address existing uncertainties over the identification of different rock types and their variations. Additionally, it holds potential to assess questions pertaining to the differentiation of the numerous transitional varieties between granitoid rocks that have been reported at the quarries of Aswan (El-Shazly 1954; Attia 1955; Higazy and Wasfy 1956) and to explore the extent to which and to establish the periods during which these different varieties of Aswan rocks were used for the production of sculpture. In order to assess the full potential of the tracing capabilities of HH-XRF methodology, however, additional geochemical data needs to be collected.
3. The results of the geochemical analyses were compared with published data for granitoids from Aswan and several granodiorites from the Eastern Desert. Aswan emerged

as most likely source for the studied rocks on the basis of consistent macroscopic and chemical observations. A connection to the exploitation of other outcrops outside of Aswan is highly unlikely.

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