

A spectral model for transient heat flow in a double U-tube geothermal heat pump system

Al-Khoury, Rafid; Focaccia, Sara

DOI

[10.1016/j.renene.2015.06.031](https://doi.org/10.1016/j.renene.2015.06.031)

Publication date

2016

Document Version

Accepted author manuscript

Published in

Renewable Energy

Citation (APA)

Al-Khoury, R., & Focaccia, S. (2016). A spectral model for transient heat flow in a double U-tube geothermal heat pump system. *Renewable Energy*, *85*, 195-205. <https://doi.org/10.1016/j.renene.2015.06.031>

Important note

To cite this publication, please use the final published version (if applicable).
Please check the document version above.

Copyright

Other than for strictly personal use, it is not permitted to download, forward or distribute the text or part of it, without the consent of the author(s) and/or copyright holder(s), unless the work is under an open content license such as Creative Commons.

Takedown policy

Please contact us and provide details if you believe this document breaches copyrights.
We will remove access to the work immediately and investigate your claim.

A Spectral Model for Transient Heat Flow in a Double U-tube Geothermal Heat Pump System

Rafid Al-Khoury¹ and Sara Focaccia²

¹Faculty of Civil Engineering and Geosciences, Computational Mechanics, Delft University of Technology,
P.O. Box 5048, 2600 GA Delft, The Netherlands

²CERENA- Centre for Natural Resources and the Environment, Instituto Superior Técnico, Lisbon
Technical University, Av. Rovisco Pais 1, 1049-001 Lisbon, Portugal

Abstract

This paper introduces a semi-analytical model based on the spectral analysis method for the simulation of transient conductive-convective heat flow in an axisymmetric shallow geothermal system consisting of a double U-tube borehole heat exchanger embedded in a soil mass. The proposed model combines the exactness of the analytical methods with an important extent of generality in describing the geometry and boundary conditions of the numerical methods. It calculates the temperature distribution in all involved borehole heat exchanger components and the surrounding soil mass using the fast Fourier transform, for the time domain; and the complex Fourier and Fourier-Bessel series, for the spatial domain. Numerical examples illustrating the model capability to reconstruct thermal response test data together with parametric analysis are given. The CPU time for calculating temperature distributions in all involved components, pipe-in, pipe-out, grout, and soil, using 16,384 FFT samples, for the time domain, and 100 Fourier-Bessel series samples, for the spatial domain, was in the order of 3 seconds in a normal PC. The model can be utilized for forward calculations of heat flow in a double U-tube geothermal heat pump system, and can be included in inverse calculations for parameter identification of shallow geothermal systems.

Keywords: Borehole heat exchanger, GSHP, TRT, spectral analysis, FFT.

1. Introduction

Geothermal heat pump (GHP) is an important source of energy for heating and cooling of buildings. It saves energy by making use of the relatively constant temperature conditions at small depths of the earth. This system, also known as borehole heat exchanger (BHE) or ground source heat pump (GSHP), works by circulating a fluid (refrigerant), mostly water with antifreeze solution, through a closed loop of polyethylene pipe that is inserted in a borehole in a soil mass. The borehole is filled with some grouted materials to fix the polyethylene pipe and to ensure a good thermal interaction with the soil. Several types of GHP are available in practice. In this publication, the GHP system is assumed to consist of a vertical double U-tube BHE embedded in a soil mass and subjected to an inlet temperature coming from the heat pump, air temperature, and a temperature coming from the bottom of the earth.

The borehole heat exchanger is a slender heat pipe with dimensions of the order of 30 mm in diameter for the U-tube, and 150 mm in diameter and 100 m in length for the borehole. The U-tube carries a working (circulating) fluid that collects heat from the surrounding soil via convection-conduction heat flow mechanisms. Physically, the heat flow process in such a system is well understood, but computationally, and in spite of the bulk of existing models, still creeping due to the combination of the slenderness of the boreholes heat exchangers and the involved thermal convection. This combination constitutes the main source of computational challenges in this field. Consequently, several theoretical and computational assumptions and approximations have been introduced in order to circumvent this problem and obtain feasible solutions. All

1 known solution techniques, such as analytical, semi-analytical and numerical, have been utilized
2 for this purpose. However, in spite of the versatility of the numerical methods, analytical and
3 semi-analytical solutions are yet preferable because of their comparatively little demands on
4 computational power and ease of use in engineering practice. In this publication, focus is placed
5 on analytical and semi-analytical solution techniques.

6 In the last three decades, several analytical and semi-analytical models for the simulation of heat
7 flow in geothermal heat pump systems with different complexities and rigor have been introduced.
8 Based on their treatment of heat flow inside the U-tubes, these models can be classified into three
9 categories: 1. No heat convection; 2. Implicit convection; and 3. Explicit convection.

10 Models belonging to the first category are those based on the work of Carslaw and Jaeger [1],
11 who seem to be the first to introduce a comprehensive treatment of heat conduction in solids.
12 Heat flow in finite, semi-infinite and infinite domains subjected to point, line, plane and
13 cylindrical heat sources were extensively studied in their work between 1947 and 1959. In the
14 meanwhile, and on the basis of Carslaw and Jaeger work, Ingersoll et al. [2] made a significant
15 contribution to the field of heat conduction in solids and provided a practical framework for
16 modeling geothermal systems. Currently, most of the analytical and semi-analytical models for
17 heat flow in geothermal heat pumps are based on these two sources. These models calculate heat
18 flow in a soil mass subjected to a heat source, representing the borehole heat exchanger,
19 regardless of the convective heat flow in the fluid inside the U-tubes and the thermal resistance
20 between the different components. Philippe et al. [3] gave a perceptive review of these models
21 and the researchers who employed them.

22 Along the same category, but different representation of the geometry, there are several other
23 models in use. In such models, the convective-conductive heat flow in the U-tubes is replaced by
24 a constant cylindrical heat source, and the geometry is described by a concentric two-dimensional
25 (radial) composite domain. Gu and O'Neal [4] gave an elaborate literature review on analytical
26 solutions of radial heat conduction in a composite domain. They utilized this technique to
27 simulate transient heat flow due to a constant heat source, resembling U-tubes, surrounded by a
28 backfill (grout) and a soil mass bounded by a far field boundary. The cross sectional areas of the
29 two branches of the U-tubes are replaced by an equivalent cross sectional area. They utilized the
30 eigenfunction expansion to solve the governing partial differential equation that gave rise to
31 solving an eight degree transcendental equation for determining the involved eigenvalues.
32 Apparently, solving an eight degree transcendental equation is difficult and might be a source of
33 numerical oscillations and computational inefficiency. In this model, summing up to 1000 terms
34 was needed for the series to converge.

35 Based on Gu and O'Neal's approach, a number of models have been introduced using different
36 mathematical formulations and solution techniques. Lamarche and Beauchamp [5] solved Gu and
37 O'Neal's composite problem using Laplace transform. They solved both forward and inverse
38 Laplace transforms analytically. Bandyopadhyay et al. [6] solved the same problem using
39 dimensionless equations by means of Laplace transform. They utilized Gaver-Stehfest numerical
40 algorithm to solve the inverse Laplace transform. Such models, together with those employing the
41 finite, infinite and cylindrical line sources, can also be classified as a no thermal resistance
42 models.

43 Models belonging to the second category are those which calculate the BHE fluid temperature
44 implicitly, i.e. without really simulating fluid flow along the axial axis of the U-tubes. In such
45 models, a mean fluid temperature is specified to indicate the average temperature in the U-tubes.
46 It is calculated by first computing the soil temperature at the borehole wall, using any of the
47 known analytical models, then adjusting the borehole thermal effective resistance to obtain
48 equilibrium. Marcotte and Pasquier [7] introduced such a model for a transient pseudo convective

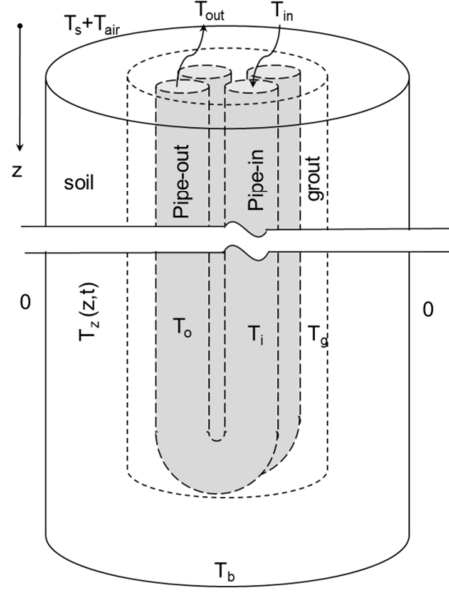
1 problem using the fast Fourier transform for discretizing the time domain, and the cubic spline for
2 interpolating results obtained at selected samples of the analytical function. They utilized the
3 principle of superposition method to simulate the response to multiple heat fluxes. Javed and
4 Claesson [8] solved Gu and O'Neal's problem using a similar pseudo convective approach.

5 Yet another type of models has also been introduced that implicitly accounts for the fluid heat
6 flow in the U-tubes. In this kind of models, heat flow in a geothermal heat pump is described by
7 an assembly of interconnected resistances and capacitors. De Carli et al. [9] and Zarrella et al. [10]
8 proposed what is known as the Capacity Resistance Model (CaRM) for the calculation of
9 transient temperature distributions in borehole heat exchangers, including those for the grout and
10 the circulating fluid. In this model, the geometry is discretized by nodes representing slices in the
11 vertical and radial directions. Heat flow in a slice is described by calculating the temperature
12 difference between adjacent slices, controlled by the thermal resistance between them. Bauer et al.
13 [11] extended the idea of the CaRM model by dividing the grout thermal resistance over the
14 number of the involved U-tubes in the borehole. Their model is known as the Thermal Resistance
15 Capacity Model (TRCM). Pasquier and Marcotte [12] extended Bauer et al. [11] model by
16 incorporating the circulating fluid and the pipe thermal capacity. They also introduced a better
17 account for the pipe spacing. This kind of models, and in spite of their apparent ease of
18 formulation, is sensitive to the number of nodes utilized to discretize the geometry, making them
19 sensitive to the thermal parameters, the definition of thermal resistance and the time steps. Such
20 models can also be classified as thermal resistance models.

21 Models belonging to the third category are those which calculate the BHE fluid temperature
22 explicitly, i.e. simulating fluid flow along the axial axis of the U-tubes. Eskilson and Claesson [13]
23 introduced a semi-analytical model for ground source heat pumps that approximates heat flow in
24 the borehole heat exchangers by two interacting channels conveying a circulating fluid in the
25 vertical axis and embedded in an axisymmetric soil mass. Heat flow in the channels is assumed
26 steady state convective, and in the soil, transient conductive. They utilized Laplace transform to
27 solve the involved heat equations of the channels, and the explicit forward difference method to
28 solve the heat equations of the soil mass. Zeng et al. [14] solved the same problem but using
29 dimensionless heat equations for the channels. This kind of models, in spite of its realistic
30 physical representation of heat flow in the GHP system, is mainly suitable for long term analyses.
31 As for the second category models, this kind of models can also be classified as thermal
32 resistance models.

33 Alongside this category, Al-Khoury [15,16] introduced a semi-analytical model for transient
34 conductive-convective heat flow in a single U-tube borehole heat exchanger embedded in a soil
35 mass. The model calculates the temperature distribution in all involved borehole heat exchanger
36 components (pipe-in, pipe-out and grout), and the surrounding soil mass using eigenfunction
37 expansion in terms of the spectral analysis method. The fast Fourier transform is utilized for
38 discretizing the time domain, and the complex Fourier series and Fourier-Bessel series are
39 utilized for discretizing the spatial domain. The main advantage of this model is that it solves the
40 governing partial differential equations of the system directly, making it physically sound.
41 Additionally, the use of the spectral analysis makes it computationally efficient.

42 In this paper, this model is extended to describe heat flow in a double U-tube borehole heat
43 exchanger embedded in an axisymmetric soil mass. Detailed mathematical formulation of the
44 double U-tube together with eigenvalue determination and spectral analysis are given. The
45 mathematical formulation and solution of heat flow in the soil mass are adopted from Al-Khoury
46 [16] but, for completeness, the general solution is described in this paper. The proposed model is
47 utilized to simulate a thermal response test (TRT).



1

2

Figure 1. A schematic representation of the double U-tube BHE temperature.

3

2. Heat flow in the double U-tube borehole heat exchanger

4

Consider a double U-tube BHE, consisting of five components (two pipes-in, denoted as $i1$ and $i2$; two pipes-out, denoted as $o1$ and $o2$; and grout, denoted as g), see Figure 1. Due to the slenderness of the BHE, the heat flow is considered to occur only along its axial axis. Radial distribution of temperature is in effect negligible. Though, there is heat exchange across the surface areas of the BHE components. Accordingly, the transient heat flow in the BHE

9

Pipes-in

10

$$\rho c_r \frac{\partial T_{i1}}{\partial t} dV_{i1} - \lambda_r \frac{\partial^2 T_{i1}}{\partial z^2} dV_{i1} + \rho c_r u \frac{\partial T_{i1}}{\partial z} dV_{i1} = b_{ig1} (T_{i1} - T_g) dS_{i1} \quad (1)$$

11

$$\rho c_r \frac{\partial T_{i2}}{\partial t} dV_{i2} - \lambda_r \frac{\partial^2 T_{i2}}{\partial z^2} dV_{i2} + \rho c_r u \frac{\partial T_{i2}}{\partial z} dV_{i2} = b_{ig2} (T_{i2} - T_g) dS_{i2}$$

12

13

Pipes-out

14

$$\rho c_r \frac{\partial T_{o1}}{\partial t} dV_{o1} - \lambda_r \frac{\partial^2 T_{o1}}{\partial z^2} dV_{o1} - \rho c_r u \frac{\partial T_{o1}}{\partial z} dV_{o1} = b_{og1} (T_{o1} - T_g) dS_{o1} \quad (2)$$

15

Grout

$$\begin{aligned}
1 \quad \rho_g c_g \frac{\partial T_g}{\partial t} dV_g - \lambda_g \frac{\partial^2 T_g}{\partial z^2} dV_g = & b_{ig1}(T_g - T_{i1})dS_{i1g} + b_{ig2}(T_g - T_{i2})dS_{i2g} + \\
& b_{og1}(T_g - T_{o1})dS_{o1g} + b_{og2}(T_g - T_{o2})dS_{o2g}
\end{aligned} \quad (3)$$

2 in which the subscripts r and g represent the circulating fluid (refrigerant) and the grout,
3 respectively; T_i , T_o and T_g are the temperatures at pipe-in, pipe-out and grout, respectively; λ_r and
4 λ_g are the thermal conductivity of the circulating fluid and grout, respectively; u (m/s) is the
5 circulating fluid velocity; b_{ig} (W/m².K) is the reciprocal of the thermal resistance between pipe-in
6 and grout; b_{og} (W/m².K) is the reciprocal of the thermal resistance between pipe-out and grout;
7 and ρc (J/ m³ K) is the volume heat capacity with c (J/kg.K) the specific heat and ρ (kg/m³) the
8 mass density. dV_{i1} is the partial volume of pipe-in(1), etc. and dS_{i1} is the partial surface area of
9 pipe-in(1), etc.

10 In practice, all U-tube pipes are made of the same materials and have the same size. This entails
11 that heat flow in pipe-in(1) is similar to pipe-in(2) and pipe-out(1) is similar to pipe-out(2),
12 leading to a reduced governing equations, which can be described as

13 Pipe-in

$$14 \quad \rho c \frac{\partial T_i}{\partial t} dV_i - \lambda \frac{\partial^2 T_i}{\partial z^2} dV_i + \rho c u \frac{\partial T_i}{\partial z} dV_i = b_{ig}(T_i - T_g)dS_i \quad (4)$$

15

16 Pipe-out

$$17 \quad \rho c \frac{\partial T_o}{\partial t} dV_o - \lambda \frac{\partial^2 T_o}{\partial z^2} dV_o - \rho c u \frac{\partial T_o}{\partial z} dV_o = b_{og}(T_o - T_g)dS_o \quad (5)$$

18 Grout

$$19 \quad \rho_g c_g \frac{\partial T_g}{\partial t} dV_g - \lambda_g \frac{\partial^2 T_g}{\partial z^2} dV_g = 2b_{ig}(T_g - T_i)dS_{ig} + 2b_{og}(T_g - T_o)dS_{og} \quad (6)$$

20 where the subscript r has been removed for clarity of notation. Note that the grout in Eq. 6 is in
21 contact with two pipes-in and two pipes-out.

22 The associated initial and boundary conditions are typically:

$$\begin{aligned}
& T_i(z,0) = T_o(z,0) = T_g(z,0) = T_s(z,0) \\
& T_i(0,t) = T_{in}(t) \\
23 \quad & T_i(L,t) = T_o(L,t) \\
& -\lambda_g \frac{\partial T_g(z,t)}{\partial z} dS_{gs} - 2b_{ig}(T_g - T_i)dS_{ig} - 2b_{og}(T_g - T_o)dS_{og} = b_{gs}(T_g - T_z)dS_{gs}
\end{aligned} \quad (7)$$

24 where, initially, the temperature distribution in the BHE components is equal to that of the steady
25 state condition of the soil before heating/cooling operation start. T_{in} is the fluid temperature at $z =$
26 0 , coming from the heat pump; T_z is the soil temperature immediately surrounding the BHE; and
27 b_{gs} is the reciprocal of the thermal resistance between the soil and the grout. dS_{ig} , dS_{og} and dS_{gs} are
28 the partial surface areas at the contact between pipe-in and grout, pipe-out and grout, and grout
29 and soil, respectively.

1 2.1 Spectral analysis

2 Integral transform methods are central among many currently applied exact solution techniques
 3 for solving transient initial and boundary value problems. The Laplace transform is one of the
 4 best known and most widely used integral transform technique. It is utilized to produce an easily
 5 solvable ordinary differential equation from a partial differential equation by transforming it from
 6 a certain domain, usually time, to the Laplace domain. However, in most cases, finding the
 7 inverse transform, which is needed to reconstruct the time function back from the Laplace domain,
 8 is quite difficult, if possible, and usually numerical and asymptotic schemes are employed in
 9 order to extract usable solutions.

10 The spectral analysis method, on the other hand, is an important alternative to the Laplace
 11 transform for solving many transient problems [18]. It is commonly utilized to transform partial
 12 differential equations in time domain to ordinary differential equations in frequency domain and
 13 vice versa. Spectral analysis of a space-time function entails discretizing the dependent variables
 14 in the frequency domain using the well-known fast Fourier transform algorithm (FFT) and
 15 discretization in the spatial domain using Fourier series expansion. It involves solving a
 16 homogeneous eigenfunction of the system to obtain its eigenvalues. The general solution of the
 17 system can then be obtained economically by summing over all significant eigenvalues, to
 18 reconstruct the spatial distribution, and the inverse fast Fourier transform algorithm (IFFT), to
 19 reconstruct the temporal distribution.

20 Using the discrete Fourier transform, a temperature function of time can be discretized as

$$21 \quad T(z, t_m) = \sum_n \hat{T}(z, \omega_n) e^{i\omega_n t_m}, \quad \hat{T}(z, \omega_n) = \frac{1}{N} \sum_m T(z, t_m) e^{-i\omega_n t_m} \quad (8)$$

22 in which N is the number of the discrete samples, where, in the fast Fourier transform, it is usually
 23 made $N = 2^\gamma = 2, 4, 8, \dots, 2048, \dots$. For a real signal, such as the one treated in this work, the
 24 transform is symmetric about a middle frequency, referred to as the Nyquist frequency. This
 25 means that N real points are transformed into $N/2$ complex points. For clarity of notation, the
 26 summation, the exponential term and the subscripts are ignored and the transform is represented
 27 as $T \Leftrightarrow \hat{T}$.

28 Applying Eq. (8) to Eqs.(4)-(6), gives

$$29 \quad \begin{aligned} & i\omega \rho c \hat{T}_i dV_i - \lambda \frac{d^2 \hat{T}_i}{dz^2} dV_i + \rho c u \frac{d\hat{T}_i}{dz} dV_i - b_{ig} (\hat{T}_i - \hat{T}_g) dS_{ig} = 0 \\ & i\omega \rho c \hat{T}_o dV_o - \lambda \frac{d^2 \hat{T}_o}{dz^2} dV_o - \rho c u \frac{d\hat{T}_o}{dz} dV_o - b_{og} (\hat{T}_o - \hat{T}_g) dS_{og} = 0 \\ & i\omega \rho_g c_g \hat{T}_g dV_g - \lambda_g \frac{d^2 \hat{T}_g}{dz^2} dV_g - 2b_{ig} (\hat{T}_g - \hat{T}_i) dS_{ig} - 2b_{og} (\hat{T}_g - \hat{T}_o) dS_{og} = 0 \end{aligned} \quad (9)$$

30 where dV_i , dV_o and dV_g are the partial volumes of pipe-in, pipe-out and grout, respectively. In this
 31 equation, the spectral representation of the time derivative has been replaced by

$$32 \quad \frac{\partial T}{\partial t} = \frac{\partial}{\partial t} \sum \hat{T}_n e^{i\omega_n t} = \sum i\omega_n \hat{T}_n e^{i\omega_n t} \Rightarrow i\omega \hat{T} \quad (10)$$

33 and the spatial derivative is replaced by

$$\frac{\partial^m T}{\partial z^m} = \frac{\partial^m}{\partial z^m} \sum \hat{T}_n e^{i\omega_n t} = \sum \frac{\partial^m \hat{T}_n}{\partial z^m} e^{i\omega_n t} \Rightarrow \frac{\partial^m \hat{T}}{\partial z^m} \quad (11)$$

2 The utilization of the spectral approach has reduced the partial differential equations, Eqs. (4)-(6),
3 to ordinary differential equations by converting the time derivative to an algebraic expression.
4 However, the resulting equations are frequency dependent and need to be solved for every
5 frequency ω_n .

6 Collecting terms, Eq. (9) can be written as

$$\begin{aligned} & -\lambda \frac{d^2 \hat{T}_i}{dz^2} dV_i + \rho c u \frac{d\hat{T}_i}{dz} dV_i + (i\omega \rho c dV_i - b_{ig} dS_{ig}) \hat{T}_i = -b_{ig} \hat{T}_g dS_{ig} \\ & -\lambda \frac{d^2 \hat{T}_o}{dz^2} dV_o - \rho c u \frac{d\hat{T}_o}{dz} dV_o + (i\omega \rho c dV_o - b_{og} dS_{og}) \hat{T}_o = -b_{og} \hat{T}_g dS_{og} \\ & -\lambda_g \frac{d^2 \hat{T}_g}{dz^2} dV_g + (i\omega \rho_g c_g dV_g - 2b_{ig} dS_{ig} - 2b_{og} dS_{og}) \hat{T}_g = -2b_{ig} \hat{T}_i dS_{ig} - 2b_{og} \hat{T}_o dS_{og} \end{aligned} \quad (12)$$

8 The associated boundary conditions in the frequency domain are

$$\begin{aligned} & \hat{T}_i(0, \omega) = \hat{T}_{in}(\omega) \\ & \hat{T}_o(L, \omega) = \hat{T}_i(L, \omega) \\ & -\lambda_g \frac{d\hat{T}_g(z, \omega)}{dz} dS_{gs} - 2b_{ig} (\hat{T}_g - \hat{T}_i) dS_{ig} - 2b_{og} (\hat{T}_g - \hat{T}_o) dS_{og} = b_{gs} (\hat{T}_g - \hat{T}_z(z, \omega)) dS_{gs} \end{aligned} \quad (13)$$

10 **Eigenfunction expansion**

11 The solution of the primary variables can be represented by an exponential complex function of
12 the form [19]:

$$\hat{T}_i = A_i e^{-ikz}, \quad \hat{T}_o = A_o e^{ikz}, \quad \hat{T}_g = A_g e^{-ikz} \quad (14)$$

14 in which A_i, A_o, A_g are the integral constants and k denotes the system eigenvalues, which need
15 to be determined. Note that different signs are employed at the exponents of Eq. (14) to impose
16 heat in pipe-in and grout to flow in $z > 0$ direction, and heat in pipe-out to flow in the opposite
17 direction, Figure 1.

18 Substituting Eq. (14) into Eq. (12), rearranging and put in a matrix form, gives

$$\begin{pmatrix} \lambda k^2 dV_i - \rho c u i k dV_i + & 0 & b_{ig} dS_{ig} \\ i\omega \rho c dV_i - b_{ig} dS_{ig} & & \\ & 2b_{ig} dS_{ig} & 2b_{og} e^{2ikz} dS_{og} \\ & & \lambda_g k^2 dV_g + i\omega \rho_g c_g dV_g - \\ & & 2b_{ig} dS_{ig} - 2b_{og} dS_{og} \\ & 0 & \lambda k^2 dV_o - \rho c u i k dV_o + \\ & & i\omega \rho c dV_o - b_{og} dS_{og} & b_{og} e^{-2ikz} dS_{og} \end{pmatrix} \begin{pmatrix} A_i \\ A_o \\ A_g \end{pmatrix} = 0 \quad (15)$$

1 Non-trivial solution of Eq. (15) can only be obtained by letting the determinate equal to zero,
 2 giving a complex six degree polynomial of the form:

$$3 \quad a_6 k^6 + a_5 k^5 + a_4 k^4 + a_3 k^3 + a_2 k^2 + a_1 k + a_0 = 0 \quad (16)$$

4 This polynomial represents the eigenfunction of the double U-tube BHE system with k denoting
 5 its set of eigenvalues, which can be obtained by solving for the roots of Eq. (16). Only for this set
 6 of eigenvalues do the eigenfunction exist that satisfy the boundary conditions of the problem. The
 7 exact forms of the coefficients of Eq. (16) are listed in the Appendix.

8 Six eigenvalues in three complex conjugates are obtained from Eq. (16), representing three basic
 9 eigenmodes, one for each BHE component. Accordingly, the solution of the temperature
 10 distribution in the three BHE components can be written as

$$\begin{aligned} \hat{T}_i &= A_i e^{-ik_1 z} + B_i e^{-ik_2 z} + C_i e^{-ik_3 z} \\ \hat{T}_g &= A_g e^{-ik_1 z} + B_g e^{-ik_2 z} + C_g e^{-ik_3 z} \\ \hat{T}_o &= A_o e^{ik_1 z} + B_o e^{ik_2 z} + C_o e^{ik_3 z} \end{aligned} \quad (17)$$

12 where the integral constants, A_i, A_o, \dots, C_g need to be determined from the boundary conditions.
 13 Since T_i, T_g , and T_o are coupled, the integral constants, A_i, A_o, \dots, C_g , are related to each other.
 14 Eq. (17) can be written as

$$15 \quad \begin{pmatrix} Q_{11} & 0 & Q_{13} \\ Q_{21} & Q_{22} & Q_{23} \\ 0 & Q_{32} & Q_{33} \end{pmatrix} \begin{pmatrix} A_i \\ A_o \\ A_g \end{pmatrix} = 0 \quad (18)$$

16 Following this equation, the relationship between the pipe-in constant and the grout constant can
 17 be expressed as

$$18 \quad A_g = A_{ig} = -\frac{Q_{11}}{Q_{13}} A_i = Y_{ig} A_i \quad (19)$$

19 Similarly, the relationship between the pipe-out constant and the grout constant can be expressed
 20 as

$$21 \quad A_g = A_{og} = -\frac{Q_{32}}{Q_{33}} A_o = Y_{og} A_o \quad (20)$$

22 For each k there is a corresponding Y_{ig} and Y_{og} , i.e. there is Y_{ig1}, Y_{og1} for k_1 , etc. (Doyle, 1988).

23 Eq. (18) states that there is a direct contact between pipe-in and the grout, and between pipe-out
 24 and the grout. There is no direct contact between the two pipes, but the grout works as the
 25 medium that transfers heat between them. The contact between pipe-in and pipe-out takes place
 26 only at the bottom of the borehole, which is not apparent in this equation. Upon solving this
 27 equations, this relatively weak coupling often leads to the generation of a spurious unphysical
 28 eigenvalue pair. This eigenvalue pair is too large compared to the other two pairs. This problem is
 29 typically encountered in solving transport phenomena using the spectral analysis method, and has
 30 been intensively treated in literature (see [20]). To obtain a solution, the spurious eigenvalues

1 have to be eliminated. One of the methods to eliminate the spurious eigenvalues is the reduction
2 of the number of the governing coupled equations.

3 Considering the geometry of the U-tube, where pipe-in meets with pipe-out at the bottom
4 boundary of the borehole, the three coupled differential equations (Eqs. (4)-(6)) can be reduced to
5 two systems of two differential equations plus an algebraic constraint linking the two zones of the
6 grout, as

7 **Pipe-in – grout**

$$\begin{aligned} \hat{T}_i &= A_i e^{-ik_1 z} + B_i e^{-ik_2 z} \\ \hat{T}_{gi} &= A_{ig} e^{-ik_1 z} + B_{ig} e^{-ik_2 z} \end{aligned} \quad (21)$$

9 **Pipe-out – grout**

$$\begin{aligned} \hat{T}_o &= A_o e^{ik_1 z} + B_o e^{ik_2 z} \\ \hat{T}_{go} &= A_{og} e^{-ik_1 z} + B_{og} e^{-ik_2 z} \end{aligned} \quad (22)$$

11 and

$$\hat{T}_g = \frac{1}{2} (\hat{T}_{gi} + \hat{T}_{go}) \quad (23)$$

13 This system is governed by the first two eigenvalues.

14 **Heat flow in Pipe-in – grout**

15 Boundary conditions relevant to Eq. (21) are:

$$\begin{aligned} \hat{T}_i(0, \omega) &= \hat{T}_{in}(\omega) \\ -\lambda_g \frac{d\hat{T}_{gi}(z, \omega)}{dz} dS_{gs} - 2b_{ig} (\hat{T}_{gi}(z, \omega) - \hat{T}_i(z, \omega)) dS_{ig} &= b_{gs} (\hat{T}_g(z, \omega) - \hat{T}_z(z, \omega)) dS_{gs} \end{aligned} \quad (24)$$

17 Substituting Eq. (21) into Eq. (24), after rearranging, gives

$$\begin{aligned} A_i + B_i &= \hat{T}_{in} \\ A_{ig} e^{-ik_1 z} (ik_1 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) + \\ B_{ig} e^{-ik_2 z} (ik_2 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) + \\ 2b_{ig} A_i e^{-ik_1 z} dS_{ig} + 2b_{ig} B_i e^{-ik_2 z} dS_{ig} &= -b_{gs} \hat{T}_z(z, \omega) dS_{gs} \end{aligned} \quad (25)$$

19 Using Eq. (19) leads, after rearrangement, to

$$\begin{aligned} A_i + B_i &= \hat{T}_{in} \\ A_i e^{-ik_1 z} [Y_{ig1} (ik_1 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) + 2b_{ig} dS_{ig}] + \\ B_i e^{-ik_2 z} [Y_{ig2} (ik_2 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) + 2b_{ig} dS_{ig}] &= -b_{gs} \hat{T}_z(z, \omega) dS_{gs} \end{aligned} \quad (26)$$

21 Putting Eq. (26) in a matrix format, and upon inverting and solving, it can be written as

$$\begin{aligned}
A_i &= \frac{1}{\Delta_i} \left[a_1 \hat{T}_{in}(\omega) - \hat{T}_z(z, \omega) \right] \\
B_i &= \frac{1}{\Delta_i} \left[a_2 \hat{T}_{in}(\omega) + \hat{T}_z(z, \omega) \right]
\end{aligned} \tag{27}$$

in which

$$\begin{aligned}
a_1 &= -\frac{e^{-ik_2z}}{b_{gs} dS_{gs}} \left[2b_{ig} dS_{ig} + Y_{ig2}(ik_2 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) \right] \\
a_2 &= \frac{e^{-ik_1z}}{b_{gs} dS_{gs}} \left[2b_{ig} dS_{ig} + Y_{ig1}(ik_1 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}) \right]
\end{aligned} \tag{28}$$

and the determinant is:

$$\begin{aligned}
\Delta_i &= \frac{e^{-ik_1z}}{b_{gs} dS_{gs}} (2b_{ig} dS_{ig} + Y_{ig1}(ik_1 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs})) - \\
&\quad \frac{e^{-ik_2z}}{b_{gs} dS_{gs}} (2b_{ig} dS_{ig} + Y_{ig2}(ik_2 \lambda_g dS_{gs} - 2b_{ig} dS_{ig} - b_{gs} dS_{gs}))
\end{aligned} \tag{29}$$

Heat flow in Pipe-out – grout

The boundary conditions relevant to Eq. (22) are:

$$\begin{aligned}
\hat{T}_o(L, \omega) &= \hat{T}_i(L, \omega) = \hat{T}_{iL} \\
-\lambda_g \frac{d\hat{T}_{go}(z, \omega)}{dz} dS_{gs} - 2b_{og} (\hat{T}_{go}(z, \omega) - \hat{T}_o(z, \omega)) dS_{og} &= b_{gs} (\hat{T}_g(z, \omega) - \hat{T}_z(z, \omega)) dS_{gs}
\end{aligned} \tag{30}$$

Substituting Eq. (22) into Eq. (30), after rearranging, gives

$$\begin{aligned}
A_o e^{ik_1L} + B_o e^{ik_2L} &= \hat{T}_{iL} \\
A_{og} e^{-ik_1z} (ik_1 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) + \\
B_{og} e^{-ik_2z} (ik_2 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) + \\
2A_o e^{ik_1z} b_{og} dS_{og} + 2B_o e^{ik_2z} b_{og} dS_{og} &= -b_{gs} dS_{gs} \hat{T}_z(z, \omega)
\end{aligned} \tag{31}$$

Using Eq. (20) leads, after rearrangement, to

$$\begin{aligned}
A_o e^{ik_1L} + B_o e^{ik_2L} &= \hat{T}_{iL} \\
A_o e^{-ik_1z} \left[2e^{2ik_1z} b_{og} dS_{og} + Y_{og1}(ik_1 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right] + \\
B_o e^{-ik_2z} \left[2e^{2ik_2z} b_{og} dS_{og} + Y_{og2}(ik_2 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right] &= -b_{gs} dS_{gs} \hat{T}_z(z, \omega)
\end{aligned} \tag{32}$$

Putting Eq. (32) in a matrix format, and upon inverting and solving, it can be written as

$$A_o = \frac{1}{\Delta_o} \left[b_1 \hat{T}_{iL} - e^{ik_2 L} \hat{T}_z(z, \omega) \right] \quad (33)$$

$$B_o = \frac{1}{\Delta_o} \left[b_2 \hat{T}_{iL} + e^{ik_1 L} \hat{T}_z(z, \omega) \right]$$

2 in which

$$b_1 = -\frac{e^{-ik_2 z}}{b_{gs} dS_{gs}} \left[2e^{2ik_2 z} b_{og} dS_{og} + Y_{og2} (ik_2 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right] \quad (34)$$

$$b_2 = \frac{e^{-ik_1 z}}{b_{gs} dS_{gs}} \left[2e^{2ik_1 z} b_{og} dS_{og} + Y_{og1} (ik_1 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right]$$

4 and the determinant is

$$\Delta_o = \frac{e^{-ik_1 z}}{b_{gs} dS_{gs}} \left[2e^{2ik_1 z} b_{og} dS_{og} + Y_{og1} (ik_1 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right] e^{ik_2 L} - \frac{e^{-ik_2 z}}{b_{gs} dS_{gs}} \left[2e^{2ik_2 z} b_{og} dS_{og} + Y_{og2} (ik_2 \lambda_g dS_{gs} - 2b_{og} dS_{og} - b_{gs} dS_{gs}) \right] e^{ik_1 L} \quad (35)$$

6 **General solution of BHE heat equations**

7 Having determined the eigenvalues and the integration constants, the general solution of the
8 double U-tube BHE system of equations can then be obtained by summing over all
9 eigenfunctions (corresponding to k_1 and k_2) and frequencies, as

10 **Pipe-in – grout**

$$T_i(z, t) = \sum_n \left(A_i e^{-ik_1 z} + B_i e^{-ik_2 z} \right) e^{i\omega_n t} \quad (36)$$

12 **Pipe-out – grout**

$$T_o(z, t) = \sum_n \left(A_o e^{ik_1 z} + B_o e^{ik_2 z} \right) e^{i\omega_n t} \quad (37)$$

14 **Grout**

$$T_g(z, t) = \frac{1}{2} \sum_n \left[\left(Y_{ig1} A_i + Y_{og1} A_o \right) e^{-ik_1 z} + \left(Y_{ig2} B_i + Y_{og2} B_o \right) e^{-ik_2 z} \right] e^{i\omega_n t} \quad (38)$$

16 where A_i and B_i are defined in Eq. (27), A_o and B_o are defined in Eq. (33), $Y_{ig1} \dots Y_{og2}$ are defined in
17 Eqs. (19)-(20) and k_1 and k_2 are determined from solving the roots of Eq. (16). The reconstruction
18 of the time domain is obtained using inverse FFT algorithm.

19 **3. Heat flow in the soil mass**

20 A detailed derivation of heat flow in the soil mass is given in Al-Khoury [16,17]. Following that,
21 the general solution of the soil heat equations in the time domain can be expressed as

$$T_{soil}(r, z, t) = T_{st}(z) + T_{tr}(z, t) + T_{rz}(r, z, t) \quad (39)$$

23 in which

$$T_{st}(z) = T_s \left(1 - \frac{z}{h}\right) + T_b \frac{z}{h} \quad (40)$$

$$T_{tr}(z, t) = \sum_n \left(-\frac{\Delta \hat{T}(\omega)}{1 - e^{-2\kappa_n h}} e^{\kappa_n(z-2h)} + \frac{\Delta \hat{T}(\omega)}{1 - e^{-2\kappa_n h}} e^{-\kappa_n z} \right) e^{i\omega_n t} \quad (41)$$

and

$$T_{rz}(r, z, t) = \sum_n \sum_m A_m J_0(\xi_m r) e^{i\omega_n t} \quad (42)$$

where

$$\begin{aligned} \kappa &= \sqrt{i\omega/\alpha} \\ \zeta &= (-\kappa^2 - \xi^2)^{1/2} \\ \xi_m &= \frac{\beta_m}{R} \end{aligned} \quad (43)$$

$$A_m = -\frac{i(e^{-i\zeta_m L} - 1)}{\zeta_m L(-i\zeta_m \lambda_s + b_{gs})} b_{gs} \hat{T}_g(z)$$

4. Reconstructing the Thermal Response Test data

The proposed spectral analysis model has been implemented in a computer code, SA-Geotherm, developed at Delft University of Technology. Both single and double U-tubes borehole heat exchangers embedded in an axisymmetric soil mass are implemented. In [15,16] full verification examples were given for heat flow in a single U-tube BHE embedded in a soil mass. Here, we examine the double U-tube BHE and its thermal interaction with the soil mass by reconstructing a real thermal response test data.

Thermal Response Test (TRT) is an in-situ parameter identification experiment for the characterization of ground thermal properties, Figure 2. It is one of the most utilized technique for determining the thermal conductivity of soil and the borehole thermal resistance.

Several analytical and numerical models have been utilized for the interpretation of TRT test results. The infinite line source model has been utilized by, among others, Mogensen [21], Hellström [22], Gehlin [23]; the finite line source model by Bandos et al. [24], to give only few examples. Numerical models have been utilized by, among others, Signorelli et al. [25], Zanchini [26], and Schiavi [27]. Nevertheless, the line source model is widely utilized for this purpose because of its simplicity of use. However, this model suffers from several shortcomings. Among others, it lumps all convective-conductive heat flow in the U-tubes and grout and their thermal interactions together with their geometry and material properties into a constant conductive infinite line source. Apparently, this simplification is a rather simplistic representation of the physics of the problem.

Here, we utilize the proposed spectral model to simulate the TRT and to back calculate the soil thermal conductivity.



Figure 2. TRT equipment

1
2

3 **TRT case study**

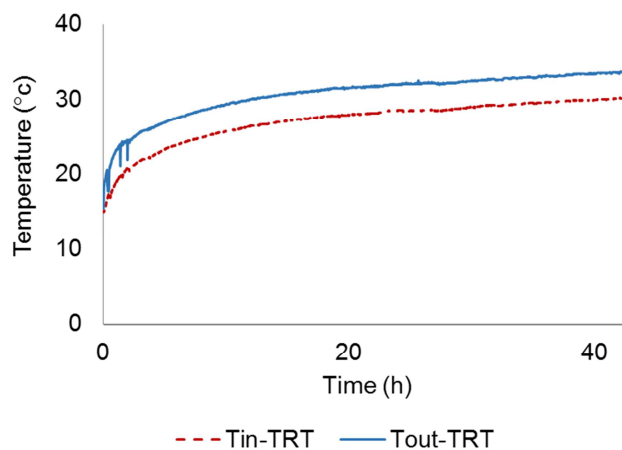
4 The thermal response test was carried out in Emilia Romagna region in Italy. The borehole heat
5 exchanger is 100 m in length and 0.127 m in diameter with a filling grout made of bentonite. The
6 collector is a double U-tube with an external diameter of 0.032 m. The working fluid is water.

7 The stratigraphy of the area is:

- 8 • 0 to 1.5 m: dry clay.
- 9 • 1.5 to 100 m: marl (there are some small infiltrations of water between 60 and 65 m depths).

10 Due to this simple stratigraphy, the soil formation surrounding the borehole was considered to be
11 consisting of a single layer. The physical and material parameters of the soil formation and the
12 TRT borehole are given in Table 1. The averaged soil thermal conductivity is not shown in the
13 table because it needs to be determined.

14 During the experiment, the fluid volume flow was measured, together with the inlet and outlet
15 temperatures. No real insulation was made to the upper part of the BHE. The experiment was
16 conducted in a cooling mode, i.e. injection of heat into the ground. Figure 3 shows the measured
17 inlet and outlet temperatures versus time.



18
19
20

Figure 3. Measured inlet and outlet temperatures during TRT.

1

Table 1. Physical and material parameters

Parameter	Value
Borehole:	
Borehole length	100 m
Borehole diameter	0.127 m
Pipe external diameter	0.032 m
Pipe thermal conductivity	0.42 W/(mK)
Fluid:	
Fluid thermal conductivity	0.56 W/(mK)
Fluid dynamic viscosity	0.001 Pa.s
Fluid velocity	0.42 m/s
Fluid specific thermal capacity	4180 J/(kg.K)
Grout:	
Grout density	1420 kg/m ³
Grout thermal conductivity	0.6 W/(m.K)
Grout specific thermal capacity	1197 J/(kg.K)
Soil:	
Soil specific thermal capacity	400 J/(kg.K)
Soil density	1680 kg/m ³

2

3 **Computational procedure**

4 In this work, no attempt was made to conduct inverse calculations by a minimization of an
5 objective function. (This will be carried out in a forthcoming work.) Rather, several spectral
6 analyses were conducted by keeping all parameters fixed, but varying the soil thermal
7 conductivity, until there is a match between the experimental results and the computed ones.

8 Time discretization of T_{in} and T_{air} signals was conducted using the forward FFT algorithm. The
9 number of samples was 16,384 (2^{14}) and the sample length was 30 s, giving a time window of
10 approximately 136 hours. Spatial discretization of the soil mass was conducted using 100 Bessel
11 function roots, and the far field boundary of the region-of-interest R was calculated as

$$12 \quad R = \sqrt{6\alpha t} \quad (44)$$

13 where α is the thermal diffusivity of the soil and t is the time when the temperature at point R
14 reaches its maximum [1]. In this work t was set equal to 100 days, giving R approximately 12 m.
15 A discussion on this choice is given in [16].

16 The thermal resistance coefficients between the borehole components and between the borehole
17 and the soil mass are calculated using the Y-configuration analogy to Ohm's law [17]. Following
18 this configuration, heat transfer coefficients for pipe-in - grout and pipe-out - grout can be described
19 as

$$20 \quad b_{ig} = \frac{1}{R_{ig}}; \quad b_{og} = \frac{1}{R_{og}} \quad (45)$$

1 where

$$2 \quad R_{ig} = R_{\text{convection}} + R_{\text{pipe material}} = \frac{1}{r_o/r_i \bar{h}} + \frac{r_o \ln(r_o/r_i)}{\lambda_p} \quad (46)$$

3 in which r_i and r_o are the inner and outer radius of pipe-in, respectively; λ_p is the thermal
4 conductivity of pipe-in material; and $\bar{h} = \text{Nu}\lambda/D$ is the convective heat transfer coefficient,
5 where D is the inner diameter of the pipe and Nu is the Nusselt Number of the circulating fluid. A
6 similar formulation is valid for R_{og} .

7 Heat transfer coefficient of the grout-soil is described as

$$8 \quad b_{gs} = \frac{1}{2R_{ig} + 2R_{og} + R_{gs}} \quad (47)$$

9 where

$$10 \quad R_{gs} = \frac{r_g \ln(r_g / r_{eq})}{\lambda_g} \quad (48)$$

11 in which r_g is the radius of the grout (borehole), and $r_{eq} = 2\sqrt{r_{in}^2 + r_{out}^2}$ with r_{in} is pipe-in inner
12 radius and r_{out} is pipe-out inner radius.

13 **Input parameters**

14 The computer code requires description of the geometry, material parameters and initial and
15 boundary conditions of both, the BHE and the soil mass. The geometry parameters include
16 information about the dimensions of the BHE components. The material parameters include
17 information about the thermal properties of the BHE components and the soil mass. The initial
18 condition includes the BHE and soil initial temperatures. The boundary conditions include the
19 input temperature at the inlet of pipe-in, together with the flow rate of the circulating fluid, and
20 the air temperature at the soil mass surface.

21 Not all information necessary to be input into the code were recorded during the experiment.
22 Though, estimates could be deduced from the available measurements. The initial temperature in
23 the soil and the borehole were assumed similar to T_{in} at the beginning, i.e. 13 °C. During the
24 experiment, the air temperature was varying between 13 and 20 °C. Table 2 shows the air
25 temperature variation with time, as was input in the code. The geometry and the material
26 parameters were input as those given in Table 1.

27

28

Table 2. Input air temperature

Time (h)	Air temperature (°C)
0	13
0.1	13.5
0.3	14
0.6	15
1.4	17
2.8	19

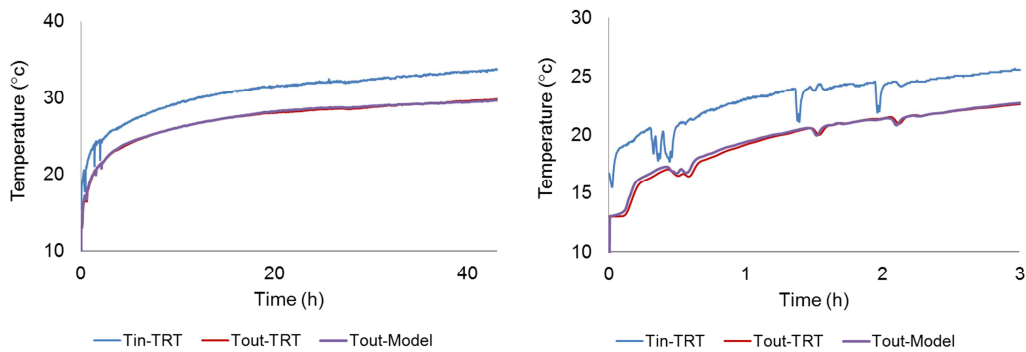
4.2	20
21	20
40	18
45	16

1 **Computational results**

2 As mentioned above, the backcalculation of the soil thermal conductivity was conducted by
 3 performing several calculations with varying soil thermal conductivity. The best fit solution is
 4 shown in Figure 4, where the soil thermal conductivity was equal to 2.15 W/m K. The figure
 5 shows the measured T_{in} and T_{out} together with the computed T_{out} .

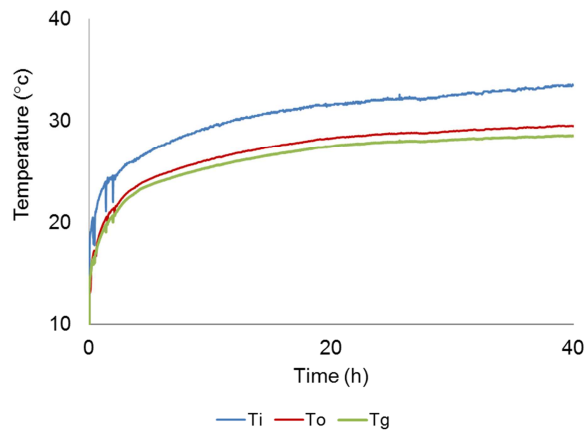
6 Apparently, the results are close and the model is capable of reconstructing the TRT in the long
 7 and short terms, as shown on the left and right hand sides of the figure respectively. An important
 8 feature of the model is manifested on the right hand side figure. The measured spikes in T_{in} data
 9 are shifted in time in the measured T_{out} data and exhibited damping. The computed T_{out} accurately
 10 exhibits these two occurrences.

11 In addition to reconstructing the TRT measured data at pipe-in inlet and pipe-out outlet, the
 12 model is capable of computing the temperature distribution at any point along the borehole and in
 13 the soil mass. For example, Figure 5 shows, in addition to pipe-in and pipe-out, the temperature
 14 distribution of the grout at the borehole surface. Figure 6 shows the temperature distribution of
 15 the soil at different radial distances from the borehole at $z = 0.25$ m.



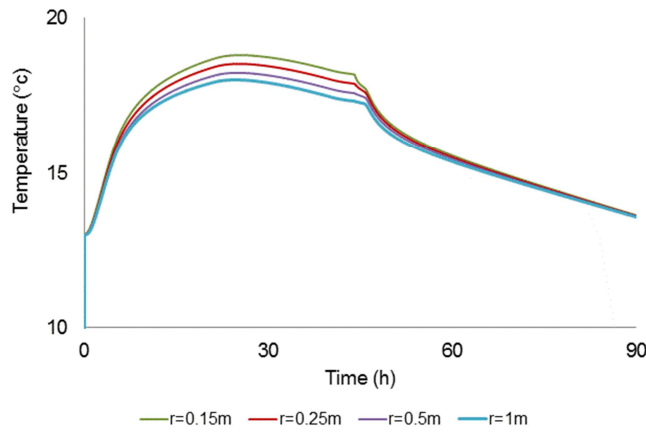
16
 17 Figure 4. Best match of the TRT data, obtained with soil thermal conductivity equal to 2.15
 18 W/mK.

19
 20



1
2

Figure 5. Temperature vs. time of pipe-in, pipe-out and grout at $z = 0$ m.



3
4
5
6

Figure 6. Soil temperature vs. time at different radial distances from the center of the borehole at depth = 0.25 m

7 **Discussion**

8 The TRT results were utilized to validate the capability of the proposed spectral model to
9 simulate heat flow in a double U-tube borehole heat exchanger and its thermal interaction with
10 the surrounding soil mass. Two outcomes can be deduced from this case study: 1. Accuracy, and
11 2. Computational efficiency.

12 For the first, the comparison between the experimental results and the computed shows that the
13 model is accurate for both, the short term and the long term. As shown in Figure 4, details of the
14 response in the short term are accurately captured by the model. In the long term, the computed
15 results are very much matching those of the experiment.

16 For the second, the proposed spectral model shares the simplicity of use of the infinite line source
17 model, but strongly overrules it in simulating the physics of the problem. This model is capable of
18 simulating full conductive-convective heat flow in all BHE components and their thermal
19 interactions between themselves and between the BHE and the soil mass. All geometrical and
20 material properties are taken into consideration.

1 The gained computational accuracy and efficiency of the proposed model make it suitable for an
2 appropriate inverse calculation based on minimization of objective functions describing the
3 difference between the experimental results and the computed. Such an inverse model is currently
4 under development.

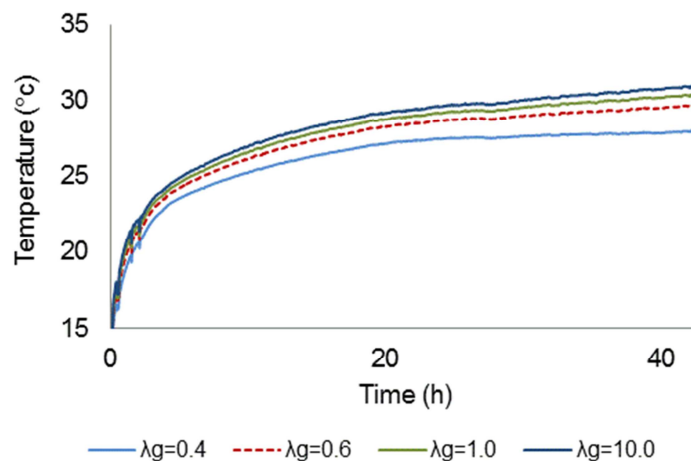
5 5. Parametric analyses

6 Having reconstructed the TRT experimental result, a parametric analysis was conducted to study
7 the effect of different material and physical parameters on heat flow in the system. Thermal
8 conductivity of the soil was kept constant. The following was studied:

9 **Grout thermal conductivity:** Figure 7 shows the temperature distribution vs. time at the outlet of
10 pipe-out for different grout thermal conductivity: $\lambda_g = 0.4, 0.6, 1.0$ and 10.0 W/m.K. The figure
11 shows that the effect of this parameter, for this set up, is not negligible.

12 This example represents a cooling mode; i.e. water with a temperature greater than that of the soil
13 is inject at the inlet of pipe-in. In spite of this, the calculated fluid temperature at the outlet is
14 higher for the high grout conductivity. This is attributed to that for a high grout conductivity,
15 more heat transfers from the inlet to the outlet via the grout; while for a low grout conductivity,
16 the opposite occurs. As heat transfer in the U-tube is mostly convective, thermal resistance
17 between the pipes and the grout is small, allowing more heat to transfer between the BHE
18 components, as compared to the conductive heat transfer between the BHE and the surrounding
19 soil. This phenomenon cannot be captured by models based on averaging the temperatures in
20 pipe-in and pipe-out and ignoring the thermal resistance and the convective heat transfer along
21 the U-tube length.

22 This issue seems in contradiction with the common practice which promotes the utilization of
23 grout with more conductivity at all times. In fact, this should not be the case. In the heating mode,
24 for instance, we tempt to gain more heat while the working fluid is running along pipe-in; and
25 preserve the gained heat while the fluid is running through pipe-out. Using high conductivity
26 grout would allow gaining more heat along pipe-in, but on the other hand, losing more heat along
27 pipe-out. Therefore, there should be an appraisal on how much heat needs to be gained and lost.
28 This should be made at the design level, depending on the initial and boundary conditions,
29 together with the required geometry and physical parameters.

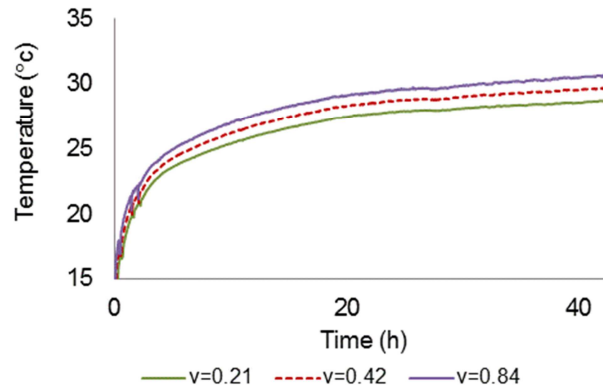


30

31

Figure 7. Temperature distribution of T_{out} for different grout thermal conductivity.

1 **Injection flow rate:** Figure 8 shows the temperature distribution vs. time at the outlet of pipe-out
 2 for different fluid velocity: $u = 0.21, 0.42$ and 0.84 m/s. The figure shows that this parameter, in
 3 the studied range, is not negligible. With higher velocity, there is a less time for the thermal
 4 interaction between the BHE and the soil, and hence higher temperature in the output. However,
 5 this is influenced by other factors, such as, the thermal conductivities of the involved components,
 6 the viscosity of the working fluid, and the geometry of the BHE. Therefore, during the design
 7 process, an appraisal between the heat flow rate and the gained temperature should be made.

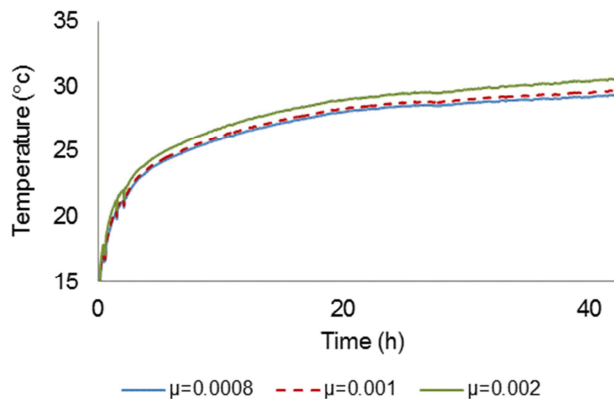


8

9 Figure 8. Temperature distribution of T_{out} for different injection flow rates.

10 **Fluid viscosity:** Figure 9 shows the temperature distribution vs. time at the outlet of pipe-out for
 11 different fluid dynamic viscosity $\mu = 0.0008, 0.001$ and 0.002 Pa.s. The fluid velocity is kept
 12 constant. The figure shows that this parameter, in the studied range, has some effect on the heat
 13 flow, but its significance would depend on the application. With higher viscosity, the circulating
 14 fluid can keep more heat, and hence brings higher temperature to the output.

15



16

17

Figure 9. Computational results of TRT by varying fluid viscosity.

18 6. Conclusions

19 A spectral model for the simulation of transient conductive-convective heat transfer in an
 20 axisymmetric shallow geothermal system consisting of a double U-tube borehole heat exchanger
 21 embedded in a soil mass is introduced. The fast Fourier transform is utilized for the discretization
 22 of the governing partial differential equations in the time domain. Complex Fourier and Fourier-

1 Bessel series are utilized for the discretization in the spatial domain. The eigenvalues of the
2 borehole heat exchanger are obtained by eigenfunction expansion, and those of the soil are
3 obtained by prescribing a homogeneous boundary condition at a fictitious boundary at some
4 distance $r = R$, where the effect of the borehole heat exchanger temperature is known a priori to
5 vanish. This condition allows for an algebraic summation over Fourier-Bessel series. This is
6 particularly important because the involved integrands are transcendental and their evaluation
7 using typical semi-infinite contour integration requires, if possible, excessive computational
8 demands.

9 The proposed model combines the exactness of the analytical methods with a great extent of
10 generality in describing the geometry and boundary conditions of the numerical methods. These
11 features make the model useful in engineering practice. The CPU time for calculating temperature
12 distributions in all involved shallow geothermal system components: pipe-in, pipe-out, grout, and
13 soil; using 16,384 FFT samples, for the time domain, and 100 Fourier-Bessel series samples, for
14 the spatial domain; was in the order of 1 second in an Intel PC.

15 As a result of the model accuracy and computationally efficiency, it can be utilized in an iterative
16 scheme for parameter identification of soil thermal parameters. In this publication, the model is
17 utilized to back calculate the soil thermal conductivity by comparing its computational results
18 with those obtained from a thermal response test. The backcalculation was conducted manually
19 by performing several calculations until a best fit is obtained. In a forthcoming work, an
20 appropriate inverse model based on minimization of the system objective function will be
21 introduced.

22 The model can be utilized for forward and inverse calculations of problems related to heat flow in
23 a double U-tube geothermal heat pump system. However, it is valid for a single layer system. For
24 a multilayer system, the model should be formulated within the spectral element method. The
25 spectral element method is a semi-analytical technique combining the exact spectral solution of
26 the system in a homogeneous domain to the finite element method solution of a heterogeneous
27 domain. The development of a spectral element model is currently underway.

28

29 **7. Appendix**

30 The coefficients of the six-degree polynomial of the double U-tube eigenfunction are:

$$31 \quad a_6 = \lambda^2 \lambda_g dV_i dV_o dV_g$$

$$32 \quad a_5 = -2i\lambda\lambda_g \rho c u dV_i dV_o dV_g$$

$$33 \quad a_4 = -b_{ig} \lambda_g \lambda dS_{ig} dV_g dV_o - \rho^2 c^2 u^2 \lambda_g dV_i dV_g dV_o + 2i\lambda\lambda_g \omega \rho c dV_i dV_g dV_o -$$

$$2\lambda^2 b_{og} dV_i dS_{og} dV_o - \lambda\lambda_g b_{og} dV_i dS_{og} dV_g + i\lambda^2 \omega \rho_g c_g dV_i dV_g dV_o -$$

$$2\lambda^2 b_{ig} dV_i dS_{ig} dV_o$$

$$34 \quad a_3 = 2\rho^2 c^2 u \lambda_g \omega dV_i dV_g dV_o + i\rho c u \lambda_g b_{og} dV_i dV_g dS_{og} + i b_{ig} \lambda_g \rho c u dV_o dV_g dS_{ig} +$$

$$4i \lambda b_{og} \rho c u dV_o dV_i dS_{og} + 2\lambda \rho c u \omega \rho_g c_g dV_o dV_i dV_g + 4i \lambda b_{ig} \rho c u dV_o dV_i dS_{ig}$$

$$\begin{aligned}
a_2 &= 2\lambda b_{ig} b_{og} dV_i dS_{ig} dS_{og} - 4i\lambda b_{ig} \omega \rho c dV_i dV_o dS_{ig} - \omega^2 \rho^2 c^2 \lambda_g dV_i dV_o dV_g - \\
&\quad ib_{ig} \omega \rho_g c_g \lambda dV_g dV_o dS_{ig} - 2\lambda \omega^2 \rho_g c_g \rho c dV_i dV_o dV_g + 2\rho^2 c^2 u^2 b_{ig} dV_i dV_o dS_{ig} - \\
1 \quad &\quad i\omega \rho c \lambda_g b_{og} dV_i dV_g dS_{og} + b_{ig} b_{og} \lambda_g dS_{ig} dV_g dS_{og} + 2\rho^2 c^2 u^2 b_{og} dV_i dV_o dS_{og} + \\
&\quad 2b_{ig} b_{og} \lambda dS_{ig} dV_o dS_{og} - 4i\lambda b_{og} \omega \rho c dV_i dV_o dS_{og} - ib_{ig} \lambda_g \omega \rho c dV_g dV_o dS_{ig} - \\
&\quad i\lambda \omega \rho_g c_g b_{og} dV_i dV_g dS_{og} - i\rho^2 c^2 u^2 \omega \rho_g c_g dV_i dV_o dV_g \\
a_1 &= 2i\rho^2 c^2 u \omega^2 \rho_g c_g dV_i dV_o dV_g - 2ib_{ig} b_{og} \rho c u dV_o dS_{ig} dS_{og} - \rho c u \omega \rho_g c_g b_{og} dV_i dV_g dS_{og} - \\
2 \quad &\quad 4\rho^2 c^2 u \omega b_{ig} dV_i dV_o dS_{ig} - 4\rho^2 c^2 u \omega b_{og} dV_i dV_o dS_{og} - b_{ig} \omega \rho_g c_g \rho c u dV_o dV_g dS_{ig} - \\
&\quad 2i\rho c u b_{ig} b_{og} dV_i dS_{ig} dS_{og} \\
a_0 &= -i\omega^3 \rho^2 c^2 \rho_g c_g dV_i dV_o dV_g + 2\omega^2 \rho^2 c^2 b_{og} dV_i dV_o dS_{og} + 2\omega^2 \rho^2 c^2 b_{ig} dV_i dV_o dS_{ig} + \\
3 \quad &\quad \omega^2 \rho c \rho_g c_g b_{og} dV_i dV_g dS_{og} + 2i\omega \rho c b_{ig} b_{og} dV_i dS_{ig} dS_{og} + 2i\omega \rho c b_{ig} b_{og} dV_o dS_{ig} dS_{og} + \\
&\quad ib_{ig} \omega \rho_g c_g b_{og} dV_g dS_{ig} dS_{og} + b_{ig} \omega^2 \rho_g c_g \rho c dV_o dV_g dS_{ig}
\end{aligned}$$

4 References

- 5 [1] Carslaw HS, Jaeger JC. Conduction of Heat in Solids. 2nd edition. Oxford University Press,
6 London, UK: 1959.
- 7 [2] Ingersoll LR, Zobel OJ, Ingersoll AC. Heat Conduction with Engineering, Geological, and
8 other Applications. Revised edition. University of Wisconsin press; 1954.
- 9 [3] Philippe M, Bernier M, Marchio D. Validity ranges of three analytical solutions to the transfer
10 in the vicinity of single boreholes. Geothermics 2009; 38, 407-413.
- 11 [4] Gu Y, O'Neal DL. An Analytical Solution to transient heat conduction in a composite region
12 with a cylindrical heat source. ASME Journal of Solar Energy Engineering 1995; 117, 242-248.
- 13 [5] Lamarche L, Beauchamp B. New solutions for the short-time analysis of geothermal vertical
14 boreholes. International Journal of Heat and Mass Transfer 2007; 50, 1408-1419.
- 15 [6] Bandyopadhyay G, Gosnold W, Mann M. Analytical and semi-analytical solutions for short-
16 time transient response of ground heat exchangers. Energy and Buildings 2008; 40, 1816-1824.
- 17 [7] Marcotte D, Pasquier P. Fast fluid and ground temperature computation for geothermal
18 ground-loop heat exchanger systems. Geothermics 2008; 37, 651-665.
- 19 [8] Javed S, Claesson J. New analytical and numerical solutions for the short-term analysis of
20 vertical ground heat exchangers. ASHRAE Transactions 2011; Vol. 117(1): 3-12.
- 21 [9] De Carli M, Tonon M, Zarrella A, Zecchin R. A computational capacity resistance model
22 (CaRM) for vertical ground coupled heat exchangers. Renewable Energy 2010; 35, 1537-1550.
- 23 [10] Zarrella A, Scarpa M, De Carli M. Short time step analysis of vertical ground-coupled heat
24 exchangers: The approach of CaRM. Renewable Energy 2011; 36, 2357-2367.
- 25 [11] Bauer P, Heidemann W, Muller-Steinhagen H, Diersch HJG. Thermal resistance and
26 capacity models for borehole heat exchangers. International Journal of Energy Research 2010;
27 Vol. 35(4), 312-320.

- 1 [12] Pasquier P, Marcotte D. Short-term simulation of ground heat exchanger with an improved
2 TRCM. *Renewable Energy* 2012; 46, 92-99.
- 3 [13] Eskilson P and Claesson J. Simulation model for thermally interacting heat extraction
4 boreholes. *Numerical Heat Transfer* 1988; 13, 149-165.
- 5 [14] Zeng H, Diao N, Fang Z. Heat transfer analysis of boreholes in vertical ground heat
6 exchangers. *International Journal of Heat and mass Transfer* 2003; 46, 4467-4481.
- 7 [15] Al-Khoury R. Spectral framework for geothermal borehole heat exchangers, *International*
8 *Journal for Numerical Methods for Heat and Fluid Flow* 2010; 20(7), 773-793.
- 9 [16] Al-Khoury R. A spectral model for shallow geothermal systems. *International Journal for*
10 *Numerical Methods for Heat and Fluid Flow* 2012; 22(1), 49-72.
- 11 [17] Al-Khoury R. Computational modeling of shallow geothermal systems. CRC Press/ Taylor
12 & Francis Group; 2012.
- 13 [18] Doyle JF. Spectral analysis of coupled thermoelastic waves. *Journal of Thermal Stresses*
14 1988; 11, 175-185.
- 15 [19] Doyle JF. *Wave Propagation in Structures: Spectral Analysis Using Fast Discrete Fourier*
16 *Transforms*. Springer-Verlag, New York; 1997.
- 17 [20] Manning ML, Bamieh B, Carlson JM. Descriptor approach for eliminating spurious
18 eigenvalues in hydrodynamic equations. Cornell University Library, arXiv:0705.1542
19 [physics.comp-ph]; 2008.
- 20 [21] Mogensen P. Fluid to Duct Wall Heat Transfer in Duct System Heat Storage. Proc. Int. Conf.
21 On Subsurface Heat Storage in Theory and Practice. Stockholm. Sweden, June 6–8, 1983, 652-
22 657.
- 23 [22] Hellström G. Thermal Analysis of Duct Storage System. Dep. of Mathematical Physics
24 University of Lund, Lund, Sweden; 1991, pp. 262.
- 25 [23] Gehlin S. Thermal response test. Method, development and evaluation. 2002; Department of
26 Environmental Engineering, University of Lulea, Sweden.
- 27 [24] Bandos T, Montero A, Fernández E, Santander JLG, Isidro JM, Pérez J et al.. Finite line-
28 source model for borehole heat exchangers: effect of vertical temperature variations. *Geothermics*
29 2009; 38, 263-370.
- 30 [25] Signorelli S, Bassetti S, Pahud D, Kohl T. Numerical evaluation of thermal response tests.
31 *Geothermics* 2007; 36, 141 – 166.
- 32 [26] Zanchini E, Terlizese T. Finite-element evaluation of thermal response tests performed on
33 U-tube borehole heat exchangers. Proceedings of the COMSOL Conference 2008, Hannover,
34 Germany.
- 35 [27] Schiavi L. 3D simulation of the Thermal Response Test in a U-tube Borehole Heat
36 Exchanger. Proceedings of the COMSOL Conference 2009, Milan.
- 37